



# UNIVERSITÀ DI PARMA

## ARCHIVIO DELLA RICERCA

University of Parma Research Repository

A dipole solution for power-law gravity currents in porous formations

This is the peer reviewed version of the following article:

*Original*

A dipole solution for power-law gravity currents in porous formations / Longo, Sandro Giovanni; Di Federico, Vittorio; Chiapponi, Luca. - In: JOURNAL OF FLUID MECHANICS. - ISSN 0022-1120. - 778:(2015), pp. 534-551. [10.1017/jfm.2015.405]

*Availability:*

This version is available at: 11381/2794671 since: 2021-10-12T19:16:18Z

*Publisher:*

Cambridge University Press

*Published*

DOI:10.1017/jfm.2015.405

*Terms of use:*

Anyone can freely access the full text of works made available as "Open Access". Works made available

*Publisher copyright*

note finali coverpage

(Article begins on next page)

13 August 2025

# A dipole solution for power-law gravity currents in porous formations

S. Longo<sup>1</sup>†, V. Di Federico<sup>2</sup>, L. Chiapponi<sup>1</sup>

<sup>1</sup>Dipartimento di Ingegneria Civile, Ambiente Territorio e Architettura (DICATeA), Università di Parma, Parco Area delle Scienze, 181/A, 43124 Parma, Italy

<sup>2</sup>Dipartimento di Ingegneria Civile, Chimica, Ambientale e dei Materiali, Università di Bologna, Viale Risorgimento, 2, 40136 Bologna, Italy

(Received ?; revised ?; accepted ?. - To be entered by editorial office)

A theoretical and experimental analysis of non-Newtonian gravity currents in porous media with variable properties is presented. A mound of a power-law fluid of flow behaviour index  $n$  is released into a semi-infinite saturated porous medium above a horizontal bed, and can drain freely out of the formation at the origin. The porous medium permeability and porosity vary along the vertical as  $z^{\omega-1}$  and  $z^{\gamma-1}$ , respectively, being  $z$  the vertical coordinate and  $\omega$  and  $\gamma$  constant numerical coefficients. A self-similar solution describing the space-time evolution of the resulting gravity current is derived for shear-thinning fluids with  $n < 1$ , generalizing earlier results for Newtonian fluids. The solution conserves a generalized dipole moment of the mound. The spreading of the current front is proportional to  $t^{\gamma n/(2+\omega(n+1))}$ . Expressions for the time evolution of outgoing flux at the origin and of the current volume are derived in closed form. The Hele-Shaw analogue is derived for flow of a power-law fluid in a porous medium with vertically variable properties. Results from laboratory experiments conducted in two Hele-Shaw cells confirm the constancy of the dipole moment and compare successfully with the theoretical formulation.

**Key words:** gravity currents, non-Newtonian, power-law, self-similar, dipole, laboratory experiments

---

## 1. Introduction

The interest in gravity currents within porous media is motivated by the need to model environmental and industrial processes in which an intruding fluid propagates inside a porous medium saturated with another fluid (Ungarish 2009).

The spreading of these currents is governed by the interplay between the pressure/buoyancy driving and viscous resistances. Semi-analytical models are available to describe their propagation in porous domains of infinite extent having plane (Huppert & Woods 1995) or radial geometry (Lyle *et al.* 2005), under instantaneous or continuous mass injection at the domain origin. A prototype case concerns the spreading of a groundwater, or liquid waste, mound injected instantaneously in a homogeneous medium of infinite extent, as it illustrates the response of the system to a pulse input; an analytical solution to this problem was derived for a constant mound volume by Barenblatt (1952) in self-similar form. In some cases the mound volume is not conserved due to a loss of fluid; this is typically the case of the response of a porous bank to a sudden surge of water level in an

† Email address for correspondence: sandro.longo@unipr.it

adjacent channel, which rapidly reverts to its original value. As a consequence, the water mound infiltrating the porous bank partially drains back into the channel. Here, while the total mass present within the porous domain decreases over time, the first spatial moment of the mound (also termed dipole moment of the flow) is conserved, as first shown in the seminal work of Barenblatt & Zel'dovich (1957). These authors derived a self-similar solution describing the spreading of the mound in the intermediate asymptotic regime, i.e. at time scales much larger than the duration of the surge initiating the motion, yet smaller than needed to reach the final equilibrium condition, or the outer boundary of the system (see Figure 1). This dipole solution has been extended to several dimensions by Hulshof & Vazquez (1993). Wagner (2005) discussed similarity solution of the second kind of the dipole problem including capillarity. Barenblatt & Vazquez (1998) extended the dipole solution to the case of assigned outgoing flux at a moving boundary. King & Woods (2003) verified the classical dipole solution performing experiments using a Hele-Shaw cell, obtaining a good agreement with the theoretical results; they then extended the model and experiments to include drainage effects. Mathunjwa & Hogg (2007) incorporated a vertical permeability variation in the dipole solution and showed it to be linearly stable.

The body of knowledge accumulated on gravity driven flows in porous media concerns mainly water and other Newtonian liquids. However, fluids flowing in natural and artificial porous media in a variety of environmental and industrial contexts often display a complex rheological nature. Relevant applications include, but are not limited to, environmental contamination and in situ remediation, enhanced oil recovery operations, filtration of polymer solutions and slurries, textiles and paper manufacturing, food processing, biological systems and biomedical techniques (e.g., Vossoughi 1999; Ciriello & Di Federico 2012, and references therein).

The nature of the non-linear relationship between stress and shear rate also varies greatly, and fluids of interest are grouped in three classes: i) time-independent, ii) time-dependent, and iii) viscoelastic. Time-independent models are the most widely used for flow problems in porous media (Wu & Pruess 1996). Among these, the power-law, or the Ostwald-de Waele model, is a two-parameter simplification of more complex relationships. It approximately describes the behaviour of time-independent non-Newtonian fluids over a certain intermediate range of shear rates and in absence of yield stress. It has been successfully adopted to describe the flow of polymer solutions, foams, and crude oils in porous media. Environmental contaminants such as slurries or dredged fluid mud (Yilmaz *et al.* 2014) and carriers for remediation agents (Tosco *et al.* 2014) are usually represented either via the power-law model, or using more complex constitutive equations including a yield stress; even in the latter case however the adoption of a pure power-law model (approximating the Herschel-Bulkley constitutive equation for vanishing values of the yield stress) permits mathematical predictions that are upper bounds for the exact solution, or may work as benchmarks for numerical codes.

The modeling of gravity currents of non-Newtonian fluids is of importance whenever density differences act as the main driving force in porous media flow of rheologically complex fluids. Most literature studies addressing non-Newtonian gravity currents in porous media adopted a power-law rheological model. The propagation of such currents was analyzed by Pascal & Pascal (1993) for constant injected volume and plane geometry, and by Bataller (2008) for assigned depth or flux at the origin and radial geometry. General solutions for a time-variable current volume were derived in two-dimensional (Di Federico *et al.* 2012*a*) and axisymmetric geometry (Di Federico *et al.* 2012*b*); the results obtained in the latter paper were confirmed by the experiments of Longo *et al.* (2013*b*).

To the best of our knowledge, the classical dipole problem (discussed above for a Newtonian fluid) was never formulated nor solved for the non-Newtonian case. This model mimics the instantaneous release of a fluid with complex rheological behaviour in a semi-infinite porous medium due to a sudden surge in level at the boundary. As such, it represents for instance the spreading of industrial effluents or slurries produced by mining activities, which are typically conveyed to disposal or tailings areas by means of open channels or flumes (Burger *et al.* 2015). An overflow in the channel causes the transported material to go over the channel walls and spread in the adjoining porous bank. Another relevant scenario is the free-surface advancement in aquifer systems of remediation agents such as nanoscale zerovalent iron (Tosco *et al.* 2014) dispersed in aqueous slurries. In both cases, the size and shape of subsurface domain interested by the spreading needs to be assessed.

In this paper, we derive a solution to the dipole problem for a non-Newtonian power-law fluid in a porous medium including the effect of spatial heterogeneity. The well-known tendency of natural media to display variations in permeability is conceptualized with a continuous power-law variation along the vertical, mimicking the sub-horizontal stratifications of porous media, or the change in fracture density with depth of densely fractured media. A vertical porosity gradient is added to the conceptual model, taking into account that in real media, variations in permeability and porosity are often interdependent (Phillips 1991; Dullien 1992). The presence of permeability and porosity gradients parallel or perpendicular to the flow direction has been shown to affect significantly the shape of gravity currents of Newtonian (Huppert & Woods 1995; Zheng *et al.* 2013) or non-Newtonian (Di Federico *et al.* 2014) fluids within porous domains.

A closed form solution to said problem is obtained in similarity form for shear thinning fluids, extending the work of King & Woods (2003), who treated the dipole flow for a Newtonian fluid and vertically variable permeability. In variance with their results, the first spatial moment is conserved only for a specific class of self-similar solutions, and the height of the fluid mound in the origin is non-zero and finite. We then explore the conditions for the validity of a Hele-Shaw analogue model, and present results from a set of laboratory experiments which test the constancy of the dipole moment, and the similarity solution obtained. In both cases a good agreement is obtained between theoretical and experimental results.

## 2. Formulation

We study the gravity-driven flow of a viscous non-Newtonian fluid having density  $\rho$  and spreading along the horizontal impermeable base of an infinite homogeneous porous domain of depth  $h_0$ , initially saturated with an ambient fluid of density  $\rho - \Delta\rho$ . The intruding fluid is described rheologically by the power-law model, which uses two parameters, flow behaviour index  $n$  and consistency index  $\tilde{\mu}$ , to express shear stress  $\tau$  as a function of shear rate  $\dot{\gamma}$  in the form  $\tau = \tilde{\mu}|\dot{\gamma}|^{n-1}\dot{\gamma}$ , valid for simple shear. The cases  $n < 1$ ,  $n = 1$ , and  $n > 1$  correspond to shear-thinning, Newtonian, and shear-thickening behaviour, respectively. For  $n = 1$ , the consistency index  $\tilde{\mu}$  becomes the dynamic viscosity  $\mu$ .

A finite mound of fluid is released in the half-plane  $x > 0$  and can drain freely at  $x = 0$ , see Figure 1. The space-time development of the current height  $h(x, t)$  is determined under the assumption of a thin gravity current having a large ratio between horizontal and vertical length scales (Bear 1972), and small height with respect to the thickness of the porous domain ( $h(x, t) \ll h_0$ ). This allows modeling the propagation as one-dimensional (neglecting vertical velocities) and disregarding the motion of the displaced

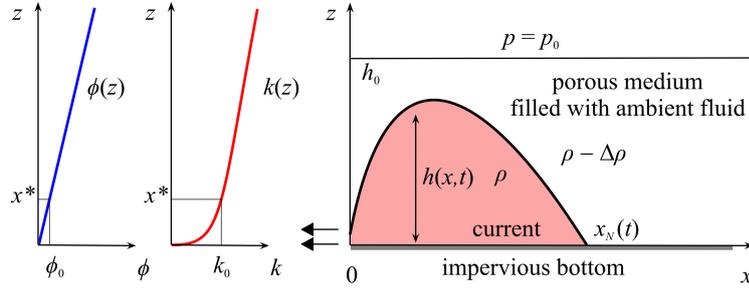


FIGURE 1. Diagram showing a mound of fluid of density  $\rho$  and height  $h(x, t)$  spreading over an impermeable base in a porous medium saturated with fluid of density  $\rho - \Delta\rho$  and draining freely at  $x = 0$ . Porosity and permeability vary with elevation  $z$  as  $\phi = \phi_0(z/x^*)^{\gamma-1}$  and  $k = k_0(z/x^*)^{\omega-1}$ . The diagrams for  $\phi$  and  $k$  refer to  $\gamma = 2$  and  $\omega = 4$ , respectively, coincident to the values adopted in our experiments.

fluid. The pressure within the intruding fluid is hydrostatic and given by

$$p(x, z, t) = p_1 + \Delta\rho gh(x, t) - \rho gz, \quad (2.1)$$

where  $p_1 = p_0 + (\rho - \Delta\rho)gh_0$  is a constant, and  $p_0$  is the constant pressure at  $z = h_0$ . The sharp interface approximation is assumed to hold; we neglect mixing due to convection assuming that no interface instabilities develop, and mixing by molecular diffusion assuming that the time scale of molecular diffusion is much longer than the time scale of the process. Surface tension effects are also neglected; this may be a limitation for very low-permeability media, as it implies disregarding the capillary entry pressure limiting the migration of the injected fluid. We further assume the validity of the modified Darcy's equation for power-law fluid flow in porous media (Pascal 1983; Shenoy 1995)

$$\mathbf{u} = -\Lambda_0^{1/n} \phi^{(n-1)/2n} k^{(n+1)/2n} |\nabla(p + \rho gz)|^{(1/n-1)} \nabla(p + \rho gz), \quad (2.2)$$

$$\Lambda_0(\tilde{\mu}, n) = \frac{1}{2\tilde{\mu}C_t} \left(\frac{50}{3}\right)^{(n+1)/2} \left(\frac{n}{3n+1}\right)^n, \quad (2.3)$$

where  $\mathbf{u}$  is the Darcy velocity,  $p$  the pressure in the intruding fluid,  $g$  the gravity,  $z$  the vertical coordinate,  $\phi$  and  $k$  the medium porosity and intrinsic permeability, respectively. Equation (2.2) derives from the capillary bundle analogy for power-law fluid flow, modified to take tortuosity effects into account via a factor  $C_t > 1$ , which has several possible expressions (for a review, see Shenoy 1995). The expression implemented here in Eq. (2.3) (for the details, see Di Federico *et al.* 2012a), is  $C_t = (25/12)^{(n+1)/2}$  (Pascal 1983). As a whole, the prefactor on the l.h.s. of (2.2) is akin to an effective fluid mobility, and depends on medium properties (porosity/permeability) in a nonlinear fashion, based on the value of the flow behaviour index  $n$ . For a Newtonian fluid,  $C_t = (25/12)$  (Christopher & Middleman 1965), the factor  $\Lambda_0$  in (2.3) reduces to  $1/\mu$ , and Equation (2.2) to Darcy's law  $\mathbf{u} = -(k/\mu)\nabla(p + \rho gz)$ .

The permeability and the porosity of the medium are assumed to vary along the vertical as

$$k = k_0 \left(\frac{z}{x^*}\right)^{\omega-1}, \quad \phi = \phi_0 \left(\frac{z}{x^*}\right)^{\gamma-1}, \quad (2.4a, b)$$

where  $x^*$  is a reference length scale,  $k_0$  and  $\phi_0$  are the permeability and porosity at  $z = x^*$ , respectively, and  $\omega$  and  $\gamma$  are constant numerical coefficients.  $\omega > 1$  ( $\gamma > 1$ )

or  $\omega < 1$  ( $\gamma < 1$ ) represent a permeability (porosity) increasing or decreasing with  $z$ , respectively. We further set  $\gamma > \gamma_1 = 0$  and  $\omega > \omega_1 = \gamma(1-n)/(n+1)$ ; this ensures the validity of the velocity scale defined in the following. The lower bounds  $\gamma_1$  and  $\omega_1$  imply that porosity and permeability do not decrease faster than  $z^{-1}$  with elevation. This constraint is of limited impact, as in most real applications to porous and fractured media the permeability tends to decrease with depth (Jiang *et al.* 2010), hence the case  $\omega > 1$  is more common than  $\omega < 1$ .

Noteworthy, in real porous media the two exponents  $\omega - 1$  and  $\gamma - 1$  are related, and their ratio typically varies between 2 and 3 (Phillips 1991; Dullien 1992), corresponding to two limiting cases of the medium microstructure, i.e. media with tubular fluid paths and with an intersecting network of cracks or fissures respectively.

For a mound of Newtonian fluid propagating in a semi-infinite homogeneous medium and draining at the origin, the volume decreases in time, but the first spatial moment  $Q = \int_0^{x_N(t)} xh(x,t)dx$ , with  $x_N$  the position of the current front, is time invariant. The constancy of  $Q$ , also termed the dipole moment, was first shown by Barenblatt & Zel'dovich (1957). The large-time behaviour of the solution is well represented by self-similar functions, which have been demonstrated to converge to the general solution of the dipole problem (see Barenblatt & Vazquez 1998). In the following, we generalize the solution of the dipole problem (King & Woods 2003) to non-Newtonian power-law fluids and vertically variable permeability and porosity, described by (2.4a,b).

Under the previous assumptions, the flow equation (2.2) simplifies to its one-dimensional version, which reads in terms of the horizontal velocity  $u(x, z, t)$  as

$$u(x, z, t) = -(\Lambda_0 \Delta \rho g)^{1/n} \phi_0^{(n-1)/(2n)} k_0^{(n+1)/(2n)} \times \left(\frac{z}{x^*}\right)^{[(\omega-1)(n+1)+(\gamma-1)(n-1)]/(2n)} \left|\frac{\partial h}{\partial x}\right|^{1/n-1} \frac{\partial h}{\partial x}, \quad (2.5)$$

taking (2.1) and (2.4a,b) into account. The local mass balance equation for a current with velocity and porosity varying along the vertical is

$$\frac{\partial}{\partial x} \left( \int_0^h u dz \right) = - \frac{\partial}{\partial t} \left( \int_0^h \phi dz \right). \quad (2.6)$$

Finally, the dipole moment  $Q$  for non homogeneous porosity is given by

$$\int_0^{x_N(t)} \int_0^{h(x,t)} x \phi dz dx = Q, \quad (2.7)$$

where  $x_N(t)$  is the horizontal extent of the current. Substituting (2.5) in (2.6-2.7) yields respectively

$$\frac{\partial h^\gamma}{\partial t} = \frac{u^*}{x^{*F_1-\gamma}} \frac{\partial}{\partial x} \left( h^{F_1} \frac{\partial h}{\partial x} \left| \frac{\partial h}{\partial x} \right|^{1/n-1} \right) \quad (2.8)$$

$$\frac{1}{x^{*\gamma+2}} \int_0^{x_N(t)} x h^\gamma dx = 1, \quad (2.9)$$

where

$$F_1 = \frac{(\omega - \gamma)(n + 1) + 2\gamma n}{2n}, \quad (2.10)$$

and

$$u^* = (\Lambda_0 \Delta \rho g)^{1/n} \frac{\gamma}{F_1} \left( \frac{k_0}{\phi_0} \right)^{(n+1)/(2n)}, \quad x^* = \left( \frac{\gamma Q}{\phi_0} \right)^{1/3} \quad (2.11)$$

are the natural velocity and length scales, respectively. The mathematical statement of the problem is completed by the boundary condition at the moving boundary

$$h(x_N(t), t) = 0. \quad (2.12)$$

As to the condition at  $x = 0$ , Barenblatt & Zel'dovich (1957) showed for a Newtonian fluid ( $n = 1$ ) that imposing the null depth at the origin is tantamount to conserving the dipole moment  $Q$ , while the volume flux draining out of the system at  $x = 0$  takes a finite nonzero value. In the following we demonstrate that for a power-law fluid, the dipole moment is conserved for a subclass of self-similar solutions of the form  $h = At^{F_3} \Phi(x/t^{F_2})$  ( $A, F_2, F_3$  being constants,  $\Phi$  a shape function), namely with  $F_3 = -2F_2/\gamma$ .

The differential problem constituted by equations (2.8) and (2.9) and boundary condition (2.12) becomes in non dimensional form

$$\frac{\partial H^\gamma}{\partial T} = \frac{\partial}{\partial X} \left[ H^{F_1} \frac{\partial H}{\partial X} \left| \frac{\partial H}{\partial X} \right|^{1/n-1} \right], \quad (2.13)$$

$$\int_0^{X_N(T)} X H^\gamma dX = 1, \quad (2.14)$$

$$H(X_N, T) = 0, \quad (2.15)$$

where the dimensionless variables are  $T = t/t^*$ ,  $X = x/x^*$ ,  $X_N = x_N/x^*$ ,  $H = h/x^*$ ,  $U = u/u^*$ , the velocity and spatial scales are defined by (2.11), and the time scale is  $t^* = x^*/u^*$ .

We observe that the balances of eqs.(2.13-2.14) require, respectively

$$\frac{H^\gamma}{T} \sim \frac{H^{F_1+1/n}}{X^{1+1/n}} \quad (2.16)$$

$$X^2 H^\gamma \sim 1. \quad (2.17)$$

Hence  $X \sim T^{F_2}$  and  $H \sim T^{F_3}$ , with

$$F_2 = \frac{\gamma n}{\omega(n+1) + 2}, \quad F_3 = \frac{-2n}{\omega(n+1) + 2}, \quad (2.18)$$

and  $F_3 = -2F_2/\gamma$ , a result independent of the nature of the fluid.

### 3. Self-similar solution

The dimensionless differential problem (2.13-2.15) admits a similarity solution of the form

$$H(X, T) = \gamma^{nF_4/(n+1)} \eta_N^{F_4} T^{F_3} \Phi(\zeta), \quad \eta = \frac{X}{T^{F_2}}, \quad \zeta = \frac{\eta}{\eta_N}, \quad (3.1)$$

where  $\eta$  is the similarity variable,  $\eta_N = \eta_N(n, \omega, \gamma)$  its value at the current front,  $\zeta$  the reduced similarity variable, and

$$F_4 = \frac{2(n+1)}{(\omega - \gamma)(n+1) + 2}. \quad (3.2)$$

Substituting (3.1) in (2.13-2.15) yields

$$\left(\Phi^{F_1} |\Phi'|^{1/n-1} \Phi'\right)' + F_2 \zeta \Phi^{\gamma-1} \Phi' - F_3 \Phi^\gamma = 0, \quad (3.3)$$

$$\eta_N = \left(\gamma^{n\gamma F_4/(n+1)} \int_0^1 \zeta \Phi^\gamma d\zeta\right)^{-1/(\gamma F_4+2)}, \quad (3.4)$$

$$\Phi(1) = 0, \quad (3.5)$$

where  $\Phi(\zeta)$  is the shape function and the prime indicates  $d/d\zeta$ . The position of the current front at any time is given by

$$X_N(T) = \eta_N T^{F_2}, \quad (3.6)$$

and its velocity is equal to

$$U_N = F_2 \eta_N T^{F_2-1}. \quad (3.7)$$

Note that the restrictions previously set on the parameter values ( $\gamma > \gamma_1$  and  $\omega > \omega_1$ ) ensure that the time exponent of the current extension  $F_2$  is positive. Correspondingly, the time exponent of the current thickness  $F_3$  is always negative, as the (time decreasing) volume of the mound is spread over a larger distance as time increases. Further, taking the partial derivatives of  $F_2$  and  $F_3$  with respect to  $n$ ,  $\omega$ , and  $\gamma$  elucidates their dependence on model parameters. Firstly, it is seen that  $F_3$  increases and  $F_2$  decreases with increasing  $\omega$ . This is so because a larger  $\omega$  implies an increase (for  $\omega > 1$ ), or a less significant decrease (for  $\omega < 1$ ), of permeability with elevation, allowing the injected mound to be thicker; correspondingly, the current extension is smaller because of mass balance. Secondly, an increase of  $\gamma$  entails a corresponding increase of  $F_2$ , as an increase (for  $\gamma > 1$ ), or a less significant decrease (for  $\gamma < 1$ ) of porosity with elevation implies that more volume is available within the porous domain, thus favouring the current spreading. Thirdly,  $F_2$  increases and  $F_3$  decreases with increasing  $n$ . The physical explanation is that the average spatial pressure gradient, proportional to time raised to  $F_3 - F_2$ , diminishes with time as  $F_3 - F_2 < 0$ . Hence the shear applied to the fluid decreases over time; this in turn implies, for a shear thinning fluid with  $n < 1$ , an increased resistance to flow, as the apparent viscosity increases with decreasing shear stress.

The flux draining into the source  $q_0(t) = h \frac{\partial h}{\partial x}|_{x=0}$ , the mound volume  $v(t) = \int_0^{x_N} \int_0^h \phi(z) dz dx$ , and the volume of fluid drained  $v_{out}(t) = \int_0^t q_0(t) dt$ , are all quantities of interest defined per unit width. Their dimensionless counterparts are equal to  $R_0 = q_0/(u^* x^*)$ ,  $V = v/x^{*2}$ , and  $V_{out} = v_{out}/x^{*2}$  respectively. After some algebra, they are derived as

$$R_0 = \phi_0 \gamma^{F_4(nF_1+1)/(n+1)-1} \eta_N^{F_4 F_1 + (F_4-1)/n} T^{-F_2-1} \left(\Phi^{F_1} \Phi' |\Phi'|^{1/n-1}\right)\Big|_{\zeta=0}, \quad (3.8)$$

$$V = \phi_0 \gamma^{\gamma n F_4/(n+1)-1} \eta_N^{\gamma F_4+1} T^{-F_2} \int_0^1 \Phi^\gamma d\zeta, \quad (3.9)$$

$$V_{out} = \frac{R_0}{F_2} T \propto T^{-F_2}, \quad (3.10)$$

where  $R_0$  and  $V_{out}$  are negative. In both expression for  $V$  and  $V_{out}$  a singularity is present for  $T \rightarrow 0$ . For  $n = 1$  (Newtonian fluid) and  $\gamma = \omega = 1$  (homogeneous medium), the governing equations reduce to  $\Phi = (\zeta^{1/2} - \zeta^2)/6$ ,  $\eta_N = 40^{1/4}$ , and  $X_N = \eta_N T^{1/4}$  (Barenblatt & Zel'dovich 1957); the dimensionless flux into the source and the mound volume per unit width become  $R_0 = -\eta_N^3 T^{-5/4}/72$  and  $V = \eta_N^3 T^{-1/4}/18$ , respectively (King & Woods 2003, with the sign of  $R_0$  changed for the different convention adopted).

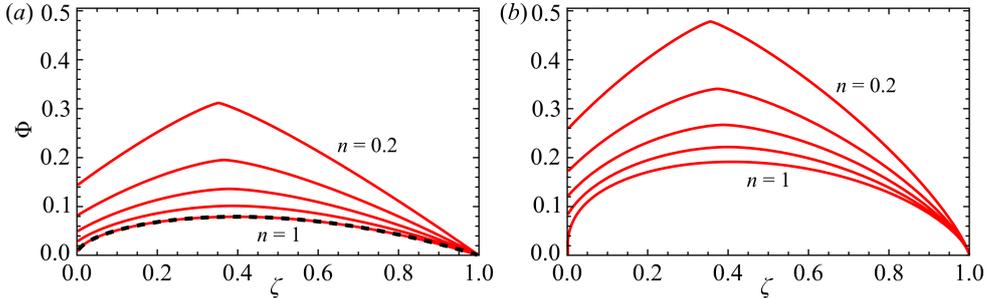


FIGURE 2. Shape function  $\Phi(\zeta)$  versus reduced similarity variable for  $n = 0.2, 0.4, 0.6, 0.8, 1$ , and (a)  $\omega = \gamma = 1$ ; (b)  $\omega = 1.75, \gamma = 1.25$ . The dots represent the analytical solution for a Newtonian fluid  $n = 1$  (Barenblatt & Zel'dovich 1957).

The numerical integration of eqs. (3.3-3.5) is conveniently performed by replacing the integral condition with an extra boundary condition on the first derivative in  $\zeta = 1$ , obtained by expanding in series the shape function near the front end.

Numerical results for the shape function  $\Phi(\zeta)$  are shown as solid curves in Figure 2 for different values of  $n$  representing shear-thinning fluids, depicting a case with uniform permeability and porosity (Figure 2a) and one with permeability and porosity increasing with  $z$  (Figure 2b). It is seen that the shape function tends to increase as  $n$  decreases; this effect is compounded when the medium properties increase with elevation. For  $n = 1$ , the numerical solution is coincident with the analytical one, which exhibits a null height at the origin. For  $n < 1$ , the shape function at  $x = 0$  is non-zero and finite. This behaviour can be interpreted as the intermediate asymptotic of a process depending on  $n, \omega, \gamma$  and having a virtual origin at negative values of  $\zeta$ ; the solution evolves from a Dirac pulse at the initial stage with a left front end starting from zero and receding to meet the virtual origin. From that moment the nonzero height condition at  $\zeta = 0$  applies (see Barenblatt & Vazquez 1998). The value of the shape function in the origin  $\Phi(0)$  for shear-thinning fluids increases as  $n$  decreases, as illustrated in Figure 3a; the decrease is more evident for low values of  $n$ , i.e. very strongly shear-thinning fluids. Figure 3a also plots the multiplicative constant  $\eta_N$  as a function of  $n$  evaluated from the numerical solution using eq. (3.4);  $\eta_N$  tends to increase with increasing values of  $n$ . The actual behaviour of the current in terms of extension and height is best analyzed in dimensional coordinates, as the scales adopted to obtain the dimensionless solution are themselves a function of model parameters.

Figure 3b depicts the flux into the source  $R_0$  and the current volume  $V$  as a function of time for different values of  $n$ . It is seen that both the flux at the origin and the mound volume decrease as the mound spreads. The flux from the current is smaller for low values of  $n$ ; as a consequence, the mound volume and its height in the origin are larger. In both cases the available analytical solution for Newtonian fluid flow in a homogeneous medium (Barenblatt & Zel'dovich 1957) is perfectly reproduced by numerical results.

The solution for shear-thickening fluids ( $n > 1$ ) indicates a detached origin on the left for positive values of  $\zeta$ , a condition also encountered by Barenblatt & Vazquez (1998) (namely in their Problem 2, *Draining problem*). The self-similar solution would exist for a infinite time and the left origin would move to the right proportionally to  $T^{-F_2}$  with a flux expressed by eq. (3.8), which should be zero, since no drain can be provided. A preliminary test conducted with a shear-thickening fluid (cornstarch suspension), showed that there is no detachment of the left origin, hence the self-similar solution (3.1) yields physically meaningless results for  $n > 1$  and will not be considered further.

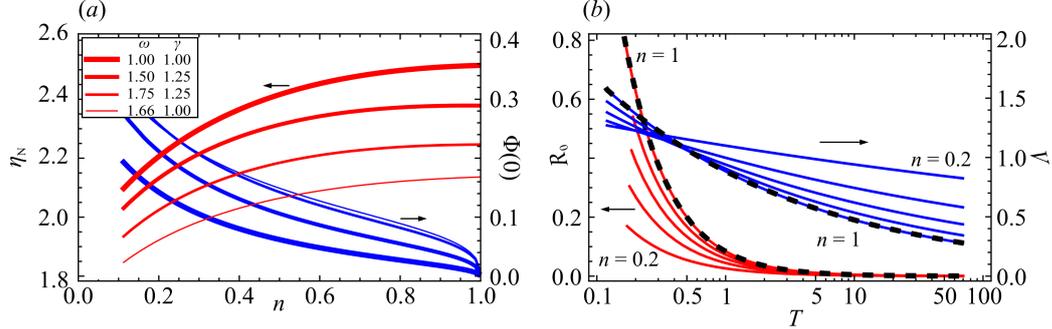


FIGURE 3. (a) Similarity variable at the current front  $\eta_N$  and shape function in the origin  $\Phi(0)$  versus  $n$  for shear-thinning power-law fluids with different conditions for porosity and permeability vertical variations. (b) The flux draining from the current  $R_0$  and the volume of the mound  $V$  as a function of time for  $n = 0.2, 0.4, 0.6, 0.8, 1$ . The dots represent the analytical solution for a Newtonian fluid (Barenblatt & Zel'dovich 1957).

#### 4. Analogue Hele-Shaw model

The Hele-Shaw approximation, a classical approach in the study of viscous flow problems, is valid for laminar flow between two plates separated by a small gap. Analogue Hele-Shaw models are traditionally used for the study of two-dimensional porous media given the identity of governing equations. Newtonian flow in a uniform Hele-Shaw cell of thickness  $b$  models Darcy flow in a homogeneous medium of permeability  $k(b) = b^2/12$ . Newtonian flow in a V-shaped Hele-Shaw cell reproduces Darcy flow in a heterogeneous medium with both permeability and porosity power-law variations along the vertical direction (Zheng *et al.* 2013); the ratio between the exponents of permeability and porosity is necessarily equal to 3, a value appropriate to represent a fissured medium. The approximation was extended to non-Newtonian power-law fluids (e.g., Aronsson & Janfalk 1992; Kondic *et al.* 1996; King 2000). Based on the similarity between the modified Darcy's law (2.2-2.3) and the laminar flow between two plates separated by a gap with half-width  $W$  varying along the vertical as  $W(z) = b_0 z^r$ , the analogy between flow of a power-law fluid in a Hele-Shaw cell of variable gap (the 'model') and a two-dimensional heterogeneous porous medium (the 'prototype') can be developed. Mapping  $H_V \rightarrow H^p$ , where  $H_V$  is the non dimensional thickness of the current in the Hele-Shaw cell and  $p > 1$  (see Appendix B for the details), the analogy requires the following correspondence amongst the parameters:

$$\begin{cases} \gamma = p(r + 1) \\ \omega = \frac{3pr(n + 1) + p(n + 3) - 2}{n + 1} \\ n = n_V, \end{cases} \quad (4.1)$$

or, in terms of  $p$  and  $r$ :

$$\begin{cases} p = \frac{(3\gamma - \omega)(n + 1) - 2}{2n} \\ r = \frac{\omega(n + 1) - \gamma(n + 3) + 2}{(3\gamma - \omega)(n + 1) - 2} \\ n = n_V, \end{cases} \quad (4.2)$$

where  $n_V$  is the behaviour index of the fluid used in the Hele-Shaw cell. It is seen that in all cases the analogy requires fluids with the same behaviour index in the prototype and in

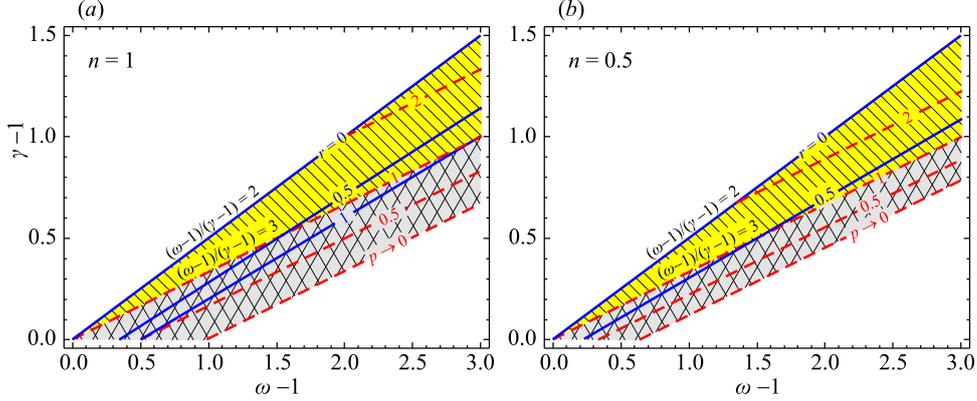


FIGURE 4. (Colour online) Interplay among parameters for Hele-Shaw analogue of flow in heterogeneous porous media with power-law permeability  $k \propto z^{\omega-1}$  and porosity  $\phi \propto z^{\gamma-1}$  variations in the vertical direction for Newtonian fluid ( $n = 1$ ); (b) for shear-thinning fluid ( $n = 0.5$ ). Dashed lines represent isovalues of  $p$ , bold lines are isovalues of  $r$ ; the isovalue  $p = 1$  is not affected by fluid behaviour index  $n$ . The closed-hatched area (yellow online) represents a positive ratio between permeability and porosity exponents in the range 2 – 3 with positive values of  $p$ ; in the closed cross-hatched area (grey online)  $p$  is also positive but  $(\omega - 1)/(\gamma - 1) > 3$ .

the model. As to the other parameters, a parallel plate Hele-Shaw cell ( $r = 0$ ) represents flow in a porous medium with a variation of properties described by  $(\omega - 1)/(\gamma - 1) = (n + 3)/(n + 1)$ . This ratio: i) tends to a value of 3, typical of fissure networks, for very strongly shear-thinning fluids with  $n \ll 1$ ; ii) is equal to 2, typical of tubular flow paths, for  $n = 1$ ; iii) is lower than 2 for shear-thickening fluids with  $n > 1$ . Mapping with  $p = 1$  can only reproduce a fixed ratio  $(\omega - 1)/(\gamma - 1) = 3$ , corresponding to fissure networks and irrespective of the fluid behaviour index  $n$ ; for example, a cell with  $r = 0.5$  reproduces the case  $\omega = 2.5$  and  $\gamma = 1.5$ . Figure 4 shows the possible combinations of parameters to ensure the analogy for fixed values of  $n$  corresponding to Newtonian and shear-thinning fluids ( $n = 0.5$ ), considering: i) positive values of  $(\omega - 1)$  and  $(\gamma - 1)$ , implying an increase of properties with elevation, the most plausible case in real porous media; ii) a positive value of the mapping parameter  $p$ , to avoid singularities at the front end, where  $H = 0$ ; iii) a shape parameter of the cell  $0 < r < 1$ , equivalent to a moderate (less than linear) increase of the gap with elevation. Inspection of the Figure 4 shows the combination of parameters satisfying the limitations  $2 \leq (\omega - 1)/(\gamma - 1) \leq 3$  and  $p > 0$ ; the former constraint is more severe.

## 5. Experiments

A series of experiments were conducted to compare with the theoretical work. The fluids were obtained by mixing glycerol, water and Xanthan gum in different ratios. In particular, 60% (vol) glycerol, 40% (vol) water, 0.1% (weight) Xanthan gum resulted in a fluid with  $n = 0.42$ ; 70% (vol) glycerol, 30% (vol) water, 0.1% (weight) Xanthan gum resulted in a fluid with  $n = 0.47$ ; 70% (vol) glycerol, 30% (vol) water, 0.06% (weight) Xanthan gum resulted in a fluid with  $n = 0.57$ . The Newtonian fluid was 95% glycerol and 5% water. The components were mixed and added with ink to enable an easy visualization of the current. The mixture was filtered in order to eliminate lumps and was stored in closed containers to avoid absorbing moisture. In general, the elimination of the lumps reduces the concentration of the Xanthan gum with respect to the nominal mass

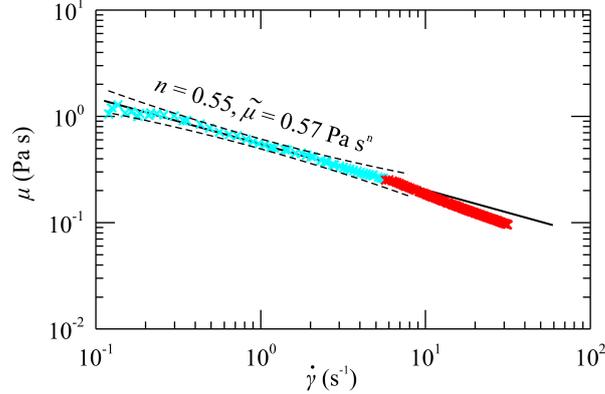


FIGURE 5. The apparent viscosity  $\mu \equiv \tilde{\mu}\dot{\gamma}^{n-1}$  as a function of shear rate; the rheometric test refers to the fluid used in Experiments 8 to 10. The bold line indicates the least squares regression for the low shear rate range, the dashed curves represent the prediction limit at 95% of confidence level.

fraction originally added. It is a source of variability in  $n$  between different mixtures with the same nominal mass fraction of Xanthan gum.

The rheological parameters were measured independently with a strain controlled parallel plate rheometer (Dynamic Shear Rheometer Anton Paar Physica MCR 101) in the low shear-rate range ( $\dot{\gamma} < 5 \text{ s}^{-1}$ ), with associated uncertainty equal to 2.9% and 3.1% for the fluid behaviour and consistency indexes, respectively. The error bounds are the prediction intervals at 95% confidence level.

Figure 5 shows the experimental data (apparent viscosity as a function of shear rate) and the fitting curve with the prediction interval for the fluid used in Experiments 8, 9, 10. A discussion on the methodology adopted for assessing the rheological parameters of the fluids is reported in Longo *et al.* (2013a). The mass density was measured by a hydrometer (STV3500/23 Salmoiraghi) with an accuracy of 0.5%.

Two different Hele-Shaw cells were used: the first had a constant gap of 3 mm or 3.2 mm; the second cell was V-shaped; the linearly varying gap reproduced a porous medium with a cubic variation of permeability and a linear variation of porosity along the vertical ( $\omega = 4, \gamma = 2$ ). Two vertical lock gates were used to confine the liquid. The lock gates were fast removed and the liquid spread along the cell and drained from the open end of the cell. The geometry of the current at different times was recorded by image analysis of frames recorded by a high-resolution video camera ( $1980 \times 1080$  pixels) recording at  $25 \text{ frames s}^{-1}$  and with a spatial resolution of  $3 \text{ pixel mm}^{-1}$ . The images were processed with a proprietary software in order to obtain a planar restitution and to extract the coordinates of the boundary between the dark current and the white background. The overall uncertainty was of  $\pm 1 \text{ mm}$ .

The parameters for all experiments are listed in Table 1. Two experiments (#2 and #7) were carried out with a Newtonian fluid to calibrate the two cells. The front position as function of the time for all experiments is shown in Figure 6a. The relevant differences between theory and experiments during the initial stage of propagation were mainly due to the finite size of the mound at  $t = 0$ ; the slumping phase is relatively short due to the fast dominance of viscous effects over inertial ones. At intermediate times, experimental data collapse onto the theoretical line. At long times the current becomes thinner and the effect of the friction at the bottom of the cell becomes non-negligible; this is reflected in a reduction of the speed of the current, with a consequent overestimation

Experiment	$\omega$	$\gamma$	$\delta$ (mm)	$Q_0^{(\gamma+2)/3}$ ( $\times 10^6 \text{ mm}^{\gamma+2}$ )	$n$	$\tilde{\mu}$ ( $\text{Pa s}^n$ )	$\rho$ ( $\text{kg m}^{-3}$ )
1	1	1	3.0	0.44	0.42	0.67	1175
2	1	1	3.0	0.44	1	1.05	1267
4	1	1	3.0	0.73	0.42	0.87	1175
6	1	1	3.2	3.10	0.42	0.87	1175
8	1	1	3.2	1.35	0.57	0.55	1174
9	1	1	3.2	0.75	0.57	0.55	1174
10	1	1	3.2	2.90	0.57	0.55	1174
7	4	2	var	3.47	1	0.90	1267
14	4	2	var	2.16	0.47	0.30	1193
16	4	2	var	1.72	0.47	0.30	1193

TABLE 1. Experimental parameters for all Experiments in Hele-Shaw cells. The last three Experiments are in a V shaped Hele-Shaw cell, reproducing flow in a porous medium with a cubic variation of permeability and a linear variation of porosity along the vertical. A power of the dipole moment (dimensions  $[L^{\gamma+2}]$ ) is listed. The nominal value of the dipole moment  $Q_0$  was computed from the initial state at rest, by digitizing the images before opening the locks.

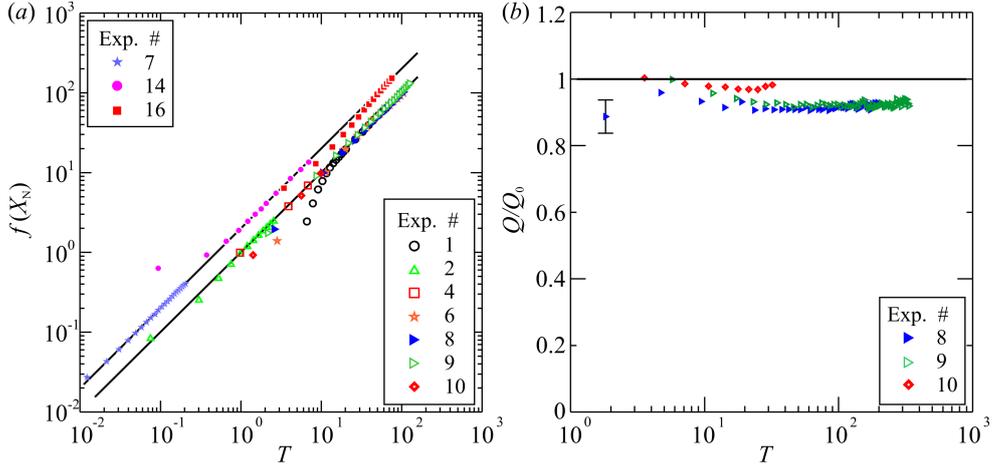


FIGURE 6. (a) Experimental front positions (symbols) and theoretical predictions (solid lines)  $X_N$  against time  $T$ . The variables  $X_N$  and  $T$  are non dimensional and  $X_N$  is scaled according to  $f(X_N) = (X_N/\eta_N)^{1/[\gamma n/(\omega(n+1)+2)]}$ . The data for experiments in the V-shaped cell are multiplied by a factor 2 in order to be separated from the data for the constant gap cell. For clarity, one point out of three is plotted. (b) Experimental dipole moment for three experiments with aspect ratio 0.8, 1.5, and 0.72, respectively.  $Q_0$  is the nominal dipole moment at time  $t = 0$ . The error bar shown indicates a single standard deviation.

of the theoretical speed. For most tests, the intermediate times are a multiple of the time scale (the time scale is generally less than 0.5 s), hence self-similarity develops after a few seconds from the opening of the locks. The long times depend on the factors determining the discrepancies between the model and the experiment. For the present tests, long times means several tens of seconds.

For a more detailed analysis of the uncertainties in the results of the model as a function of the uncertainties in the parameters, the methodology described in Di Federico *et al.* (2014) can be applied. Figure 6b shows the dipole moment as function of time for three selected experiments (#8, 9, and 10), characterised by initial aspect ratios (ratio of the

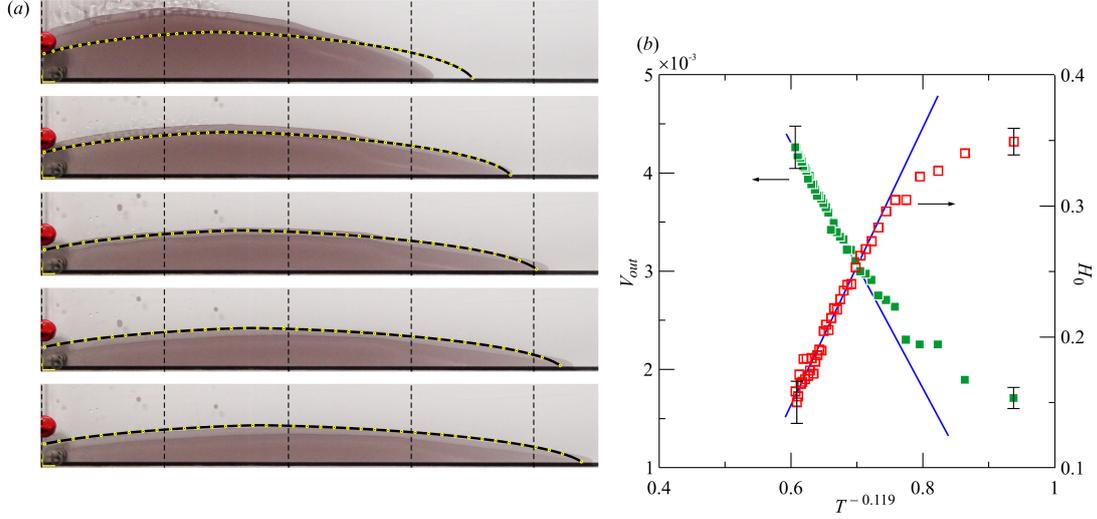


FIGURE 7. (Colour online) Experiment #16 (a) Photographs showing the current profiles at various times since release (10, 30, 50, 70 and 90 s) in a V-shaped Hele-Shaw cell. The curves with symbols represent the theoretical profile. The vertical dashed lines are 10 cm apart. (b) The dimensionless volume of fluid drained at the origin  $V_{out}$  (defined by Eq. 3.10) plotted against  $T^{-F_2} \equiv T^{-0.119}$  (filled symbols) and the dimensionless (scaled with  $x_*$ ) height of the current in the origin plotted against  $T^{F_3} \equiv T^{-F_2}$  (empty symbols). The time increases toward the origin. The error bars indicate a single standard deviation.

initial height to initial length of the fluid behind the lock gates) equal to 0.8, 1.5, and 0.72, respectively. After an initial slumping phase, the moment tends to a constant, as predicted by the theory. The moment was estimated by fitting the experimental points representing the boundary of the current (detected with the video image analysis) and then integrating the fitting function. The uncertainty in the dipole moment has been evaluated with a Monte Carlo simulation, by (i) generating several samples of points (10 000 samples) representing the boundary, by assuming a Gaussian distribution with variance equal to 1 mm around the experimental value; (ii) fitting the synthetic samples and integrating the fitting curves; (iii) computing the variance of the samples of the results of the integration. A comparison of the experimental profile with theory is shown in Figure 7a. Near the origin the height of the current is non zero, as predicted by the theory. Near the front, capillarity effects, not included in the theory, produce an underestimation of the height and of the front position. Elsewhere, the agreement between theory and experiments is remarkably good at intermediate and late times, indicating that the current converges to the self-similar regime. A balance between two counteracting effects not included in the model, i.e. the depression of the current associated with the vertical velocity and capillarity, probably contributes to the excellent fit.

Finally, we investigate the current behaviour in the origin by plotting the volume of fluid drained  $V_{out}$  and the height  $H_0$  against time for Experiment #16 (Figure 7b). Theoretically,  $V_{out}$  and  $H_0$  are proportional to  $T^{F_2}$  and  $T^{-F_3}$  respectively. In the selected case,  $F_3 = -F_2$  as  $\gamma = 2$ , hence the same exponent is expected experimentally for the two plots except for the sign. Inspection of Figure 7b confirms the theory for intermediate and late times. Similar plots are obtained for the other experiments.

## 6. Conclusions

We have analyzed the intermediate asymptotic behaviour of a mound of power-law shear-thinning fluid propagating in a two-dimensional porous medium freely draining at one end, considering constant or vertically varying porosity and permeability according to a power-law relationship. This type of variation mimics the sub-horizontal stratifications of porous media, or the change in fracture density with depth of densely fractured media. A similarity solution has been derived in closed form, extending in two respects the work of King & Woods (2003), who treated the dipole flow for a Newtonian fluid and vertically variable permeability. The dipole moment of the flow is conserved with a non zero height of the current at the origin. We have established an analogy with the flow of a fluid with the same behaviour index in a Hele-Shaw cell with half-gap varying in the vertical as  $W(z) = b_0 z^r$ , representing a variety of combinations of porosity and permeability. The general Hele-Shaw analogue provides a useful tool for the experimental simulation of power-law free-surface flows in porous media. A set of experiments based on this analogy were conducted in a constant gap ( $r = 0$ ) and V-shaped cell ( $r = 1$ ). The time series of the front-end position, dipole moment, drained volume, and height in the origin, as well as the current profiles, agreed well with the theoretical predictions at intermediate and late times. Our results demonstrate that the combined effect of fluid rheology and spatial variations of medium properties significantly affect the current extension and profile, and extend the available body of knowledge on non-Newtonian flow in porous media, a topic of interest for its environmental and industrial implications.

## Acknowledgements

We thank the anonymous Reviewers for their useful comments and suggestions that helped in improving the quality and the readability of the manuscript.

## Appendix A. The integration of the differential problem

The integration of the differential problem in (3.3–3.5) can be obtained with standard numerical algorithms (we used those implemented in Wolfram Mathematica<sup>®</sup> 10) and requires an extra condition near the front end. Here, for  $\zeta \rightarrow 1 - \epsilon$  the power series expansion of the shape function yields

$$\Phi(\epsilon) = \left( \frac{F_2}{\gamma F_5^{1/n}} \right)^{nF_5} \epsilon^{F_5}, \quad (\text{A.1})$$

$$\Phi'(\epsilon) = -F_5 \left( \frac{F_2}{\gamma F_5^{1/n}} \right)^{nF_5} \epsilon^{F_5-1}, \quad (\text{A.2})$$

where

$$F_5 = \frac{2}{(\omega - \gamma)(n + 1) + 2} \quad (\text{A.3})$$

and  $\epsilon$  is a small positive quantity ( $\epsilon = 10^{-6} - 10^{-8}$  in most computations). However, numerical instabilities develop near the maximum for  $n \neq 1$  due to the absolute value operator in (3.3). These instabilities can be overcome by increasing the accuracy and the precision in the algorithm application. For small  $n$ , we found more efficient to split the integration into a right branch from  $\zeta = 1$  to  $\zeta = \zeta_M$  ( $\zeta_M$  is the abscissa where the maximum value  $\Phi_M$  occurs), and a left branch from  $\zeta = \zeta_M$  to  $\zeta = 0$ . A series expansion

of the shape function near  $\zeta_M$ , valid for  $n < 1$  gives

$$\Phi(\zeta_M - \epsilon) = \Phi_M - \left( \frac{(1-n)^{1/n} F_2 \zeta_M}{\Phi_M^{1+F_1-\gamma}} \right)^{n/(1-n)} \zeta_M^{1/(1-n)} \epsilon^{1/(1-n)}, \quad (\text{A.4})$$

$$\Phi'(\zeta_M - \epsilon) = \frac{1}{1-n} \left( \frac{(1-n)^{1/n} F_2 \zeta_M}{\Phi_M^{1+F_1-\gamma}} \right)^{n/(1-n)} \zeta_M^{n/(1-n)} \epsilon^{n/(1-n)}. \quad (\text{A.5})$$

## Appendix B. Analogue Hele-Shaw model of power-law fluid in a porous medium

In developing the present analogy, we assume that the cell gap and the porous medium properties (permeability and porosity) are constant in the horizontal direction (parallel to flow), and vary in the vertical direction (transverse to flow). We consider the momentum and local mass balance equations for a gravity current of a power-law fluid between two surfaces with varying half-width along the vertical, given by the monomial expression  $W(x) = b_0 z^r$ . The hypothesis of no slip condition at the rigid walls leads to the following partial differential equation for the height of the current

$$\frac{\partial h_v^{r+1}}{\partial t} = \left[ \left( \frac{\Delta \rho_v g}{\tilde{\mu}_v} \right)^{1/n_v} \frac{n_v^2 (r+1)}{[r(2n_v+1) + n_v](2n_v+1)} b_0^{(n_v+1)/n_v} \right] \times \frac{\partial}{\partial x} \left[ h_v^{1+r(2n_v+1)/n_v} \frac{\partial h_v}{\partial x} \left| \frac{\partial h_v}{\partial x} \right|^{1/n_v-1} \right], \quad (\text{B.1})$$

where the subscript v refers to the Hele-Shaw cell, which is to be solved with the boundary condition at the front end  $h_v(x_N(t), t) = 0$  and respecting the integral constraint

$$\frac{2b_0}{r+1} \int_0^{x_N(t)} x h_v^{r+1} dx = Q_v, \quad (\text{B.2})$$

where  $Q_v$ , having dimensions  $[L^{3+r}]$ , is the constant dipole moment. By introducing the following length and velocity scales

$$x_v^* = \left( \frac{r+1}{2b_0} Q_v \right)^{1/(r+3)} \quad (\text{B.3})$$

$$u_v^* = \left( \frac{\Delta \rho_v g}{\tilde{\mu}_v} \right)^{1/n_v} \frac{n_v^2 (r+1) p^{1/n_v}}{[r(2n_v+1) + n_v](2n_v+1)} \left( \frac{r+1}{2} b_0^{3/r} Q_v \right)^{r(n_v+1)/[n_v(r+3)]}, \quad (\text{B.4})$$

where the coefficient  $p^{1/n_v}$  ( $p > 0$ ) is introduced for convenience, eqs.(B.1-B.2) become

$$\frac{\partial h_v^{r+1}}{\partial t} = \frac{u_v^*}{x_v^{*r(n_v+1)/n_v} p^{1/n_v}} \frac{\partial}{\partial x} \left[ h_v^{1+r(2n_v+1)/n_v} \frac{\partial h_v}{\partial x} \left| \frac{\partial h_v}{\partial x} \right|^{1/n_v-1} \right], \quad (\text{B.5})$$

$$\frac{1}{x_v^{*r+3}} \int_0^{x_N(t)} x h_v^{r+1} dx = 1. \quad (\text{B.6})$$

Equations (B.5-B.6) become in non dimensional form

$$\frac{\partial H_v^{r+1}}{\partial T} = \frac{1}{p^{1/n_v}} \frac{\partial}{\partial X} \left[ H_v^{1+r(2n_v+1)/n_v} \frac{\partial H_v}{\partial X} \left| \frac{\partial H_v}{\partial X} \right|^{1/n_v-1} \right], \quad (\text{B.7})$$

$$\int_0^{X_N(T)} X H_V^{r+1} dX = 1. \quad (\text{B.8})$$

Mapping  $H_V \rightarrow H^p$  transforms eqs. (B.7-B.8) into

$$\frac{\partial H^{p(r+1)}}{\partial T} = \frac{\partial}{\partial X} \left[ H^{p[1+r(2n_v+1)/n_v]+(p-1)/n_v} \frac{\partial H}{\partial X} \left| \frac{\partial H}{\partial X} \right|^{1/n_v-1} \right], \quad (\text{B.9})$$

$$\int_0^{X_N(T)} X H^{p(r+1)} dX = 1. \quad (\text{B.10})$$

The analogy between the model Hele-Shaw flow described by eqs. (B.9-B.10) and the prototype porous flow described by eqs. (2.13–2.14), requires the corresponding terms in the differential and integral expressions to be identical. Equating the exponents yields the solutions shown in equations (4.1) and (4.2). The established analogy allows to perform experiments in a Hele-Shaw cell and transform the experimental results obtained (the measured height of the current and the time series of the front end position) into equivalent results for the mound spreading in a porous medium.

#### REFERENCES

- ARONSSON, G. & JANFALK, U. 1992 On Hele-Shaw flow of power-law fluids. *Euro. Jnl of Appl. Math.* **3**, 343–366.
- BARENBLATT, G. I. 1952 On some unsteady motions of fluids and gases in porous medium. *Prikl. Math. Mekh.* **16**(1), 67–78.
- BARENBLATT, G. I. & VAZQUEZ, J. L. 1998 A new free boundary problem for unsteady flows in porous media. *Euro Jnl of Appl. Math.* **9**, 37–54.
- BARENBLATT, G. I. & ZEL'DOVICH, Y. 1957 On dipole solutions in problems of non-stationary filtration of gas under polytropic regime. *Prikl. Math. Mekh.* **21**(5), 718–720.
- BATALER, R.C. 2008 On unsteady gravity flows of a power-law fluid through a porous medium. *Appl. Math. Comput.* **196**, 356–362.
- BEAR, J. 1972 *Dynamics of Fluids in Porous Media*. Dover Publications.
- BURGER, J.H., HALDENWANG, R. & ALDERMAN, N.J. 2015 Laminar and turbulent flow of non-newtonian fluids in open channels for different cross-sectional shapes. *J. Hydraul. Eng.* **141**(4), 04014084.
- CIRIELLO, V. & DI FEDERICO, V. 2012 Similarity solutions for flow of non-Newtonian fluids in porous media revisited under parameter uncertainty. *Adv. Wat. Res.* **43**, 38–51.
- CRISTOPHER, R. H. & MIDDLEMAN, S. 1965 Power-law flow through a packed tube. *Ind. Eng. Chem. Fundam.* **4**, 422–427.
- DI FEDERICO, V., ARCHETTI, R. & LONGO, S. 2012a Similarity solutions for spreading of a two-dimensional non-Newtonian gravity current. *J. Non-Newtonian Fluid Mech.* **177–178**, 46–53.
- DI FEDERICO, V., ARCHETTI, R. & LONGO, S. 2012b Spreading of axisymmetric non-Newtonian power-law gravity currents in porous media. *J. Non-Newtonian Fluid Mech.* **189–190**, 31–39.
- DI FEDERICO, V., LONGO, S., CHIAPPONI, L., ARCHETTI, R. & CIRIELLO, V. 2014 Radial gravity currents in vertically graded porous media: Theory and experiments for Newtonian and power-law fluids. *Adv. Wat. Res.* **70**, 65–76.
- DULLIEN, F. A. 1992 *Porous Media: Fluid Transport and Pore Structure*. Academic Press.
- HULSHOF, J. & VAZQUEZ, J. L. 1993 The dipole solution for the porous medium equation in several space dimensions. *Annali della Scuola Normale Superiore di Pisa - Classe di Scienze - Ser. 4* **20**(2), 193–217.
- HUPPERT, H. E. & WOODS, A. W. 1995 Gravity-driven flows in porous layers. *J. Fluid Mech.* **292**, 55–69.
- JIANG, X.-W., WANG, X.-S. & WAN, L. 2010 Semi-empirical equations for the systematic

- decrease in permeability with depth in porous and fractured media. *Hyd. J.* **18(4)**, 839–850.
- KING, S. E. 2000 Non-Newtonian gravity currents in Hele-Shaw cells. MSc thesis, University of Bristol.
- KING, S. E. & WOODS, A. W. 2003 Dipole solutions for viscous gravity currents: theory and experiments. *J. Fluid Mech.* **483**, 91–109.
- KONDIC, L., PALFFY-MUHORAY, P. & SHELLEY, M. J. 1996 Models of non-Newtonian Hele-Shaw flow. *Phys. Rev. E* **54(5)**, 4536–4539.
- LONGO, S., DI FEDERICO, V., ARCHETTI, R., CHIAPPONI, L., CIRIELLO, V. & UNGARISH, M. 2013*a* On the axisymmetric spreading of non-newtonian power-law gravity currents of time-dependent volume: An experimental and theoretical investigation focused on the inference of rheological parameters. *Journal of Non-Newtonian Fluid Mechanics* **201** (0), 69 – 79.
- LONGO, S., DI FEDERICO, V., CHIAPPONI, L. & ARCHETTI, R. 2013*b* Experimental verification of power-law non-Newtonian axisymmetric porous gravity currents. *J. Fluid Mech.* **731**, **R2**, 1–12.
- LYLE, S., HUPPERT, H.E., HALLWORTH, M., BICKLE, M. & CHADWICK, A. 2005 Axisymmetric gravity currents in a porous medium. *J. Fluid Mech.* **543**, 293–302.
- MATHUNJWA, J.S. & HOGG, A.J. 2007 Freely draining gravity currents in porous media: Dipole self-similar solutions with and without capillary retention. *Euro. Jnl Appl. Math.* **18**, 337–362.
- PASCAL, H. 1983 Nonsteady flow of non-Newtonian fluids through a porous medium. *Int. J. Eng. Sci.* **21**, 199–210.
- PASCAL, J.P. & PASCAL, H. 1993 Similarity solutions to gravity flows of non-Newtonian fluids through porous media. *Int. J. Non-Linear Mech* **28(2)**, 157–167.
- PHILLIPS, O. M. 1991 *Flow and Reactions in Porous Rocks*. Cambridge University Press.
- SHENOY, A. V. 1995 Non-Newtonian fluid heat transfer in porous media. *Adv. Heat Trans.* **24**, 102–190.
- TOSCO, T., PAPINI, M., VIGGI, C. & SETHI, R. 2014 Nanoscale zerovalent iron particles for groundwater remediation: a review. *J. Clean Prod.* **77**, 10–21.
- UNGARISH, M. 2009 *An Introduction to Gravity Currents and Intrusions*. CRC Press.
- VOSSOUGH, S. 1999 *Advances in the Flow and Rheology of Non-Newtonian Fluids, Vol. 2*, chap. Flow of Non-Newtonian Fluids in Porous Media, pp. 1183–1236. Elsevier Science Publishing Company.
- WAGNER, B. 2005 An asymptotic approach to second-kind similarity solutions of the modified porous-medium equation. *J. Engrng Math.* **53**, 201–220.
- WU, Y.-S. & PRUESS, K. 1996 *Advances in Porous Media, Volume 3*, chap. Flow of non-Newtonian fluids in porous media, pp. 87–184. Elsevier.
- YILMAZ, N.A., TESTIK, F.Y. & CHOWDURY, M.R. 2014 Laminar bottom gravity currents: friction factor-Reynolds number relationship. *J. Hydr. Res.* **52(4)**, 545–558.
- ZHENG, Z., SOH, B., HUPPERT, H E & STONE, H. A. 2013 Fluid drainage from the edge of a porous reservoir. *J. Fluid Mech.* **718**, 558–568.