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## 2 **Laboratory sandbox validation of pollutant source** 3 **location methods**

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8

### 9 **Abstract**

10 Inverse methods can be used to recover the pollutant source location from concentration data. In this paper, the relative  
11 effectiveness of two proposed methods, simultaneous release function and source location identification (SRSI) and  
12 backward probability model based on adjoint state method (BPM-ASM) are evaluated using real data collected by using  
13 experimental equipment. The device is a sandbox that reproduces an unconfined aquifer in which all the variables are  
14 controlled. A numerical model was calibrated using experimental observations. The SRSI is a stochastic procedure  
15 which finds the source location and the release history by means of a Bayesian geostatistical approach. The BPM-ASM  
16 provides the backward probability location of the pollutant detected at a monitoring point by means of a reverse  
17 transport simulation. The results show that both methods perform well. While the simultaneous release function and  
18 source location identification method requires a preliminary delineation of a probable source area and some weak  
19 hypotheses about the statistical structure of the unknown release function, the backward probability model requires  
20 some hypothesis about the contaminant release time. A case study was performed using two observation points only,  
21 and despite the scarcity of data, both methodologies were able to accurately reconstruct the true source location. The  
22 geostatistical approach has the advantage to recover the release history function too, whilst the backward probability  
23 model works well with fewer data. If there are many observations, both methodologies may be computationally heavy.  
24 A transfer function approach has been adopted for the numerical definition of the sensitivity matrix in the SRSI method.  
25 The reliability of the experimental equipment was tested in previous laboratory works, conducted under several  
26 different conditions.

27 *Keywords: Geostatistical approach; Transfer function; Source detection; Backward Location PDF; Sandbox*

### 28 **List of symbols**

29  $\mathbf{b}$   $p \times 1$  unknown coefficients

30  $C(\mathbf{x}, t)$  concentration at point  $\mathbf{x}$  and time  $t$

31  $C_0$  solution concentration

32  $\hat{C}_i$  observed concentration

33  $C_w$  initial condition in backward model

34  $\mathbf{D}$  dispersion tensor

35  $F(t)$  concentration of the water injected at the source as function of time  $t$

36  $F_0$  constant and known mass rate input function

37	$f_{\hat{c}_i}$ measured concentrations PDF
38	$f(\mathbf{x},t)$ transfer function at position $\mathbf{x}$ and time $t$
39	$f_{\mathbf{x}}(\mathbf{x},t)$ backward location PDF
40	$\mathbf{H}$ $m \times n$ sensitivity matrix
41	$\mathbf{h}(\mathbf{s})$ $m \times 1$ vector that describes the transport process
42	$H_L(\mathbf{x},t)$ load term
43	$H_{SF}(t)$ Heaviside step function
44	$h_D$ head level downstream
45	$h_U$ head level upstream
46	$J$ generic sub-areas
47	$K$ hydraulic conductivity
48	$m$ number of observations
49	$m_o$ release source mass
50	$M$ random source mass
51	$\mathbf{M}$ $p \times n$ multipliers
52	$n$ number of unknowns
53	$\mathbf{n}$ normal versor
54	$p$ number of unknown coefficients
55	$\mathbf{Q}(\boldsymbol{\theta})$ $n \times n$ matrix, covariance of the unknown process
56	$Q_{in}$ injected flow rate
57	$Q_w$ initial condition in backward model
58	$q_0$ injected solution discharge
59	$q_I$ source inflow rate per unit volume
60	$\mathbf{R}$ $m \times m$ error covariance matrix
61	$\mathbf{s}$ $n \times 1$ unknowns
62	$s(t)$ unknown release function
63	$\tilde{s}$ transformed unknown function
64	$\hat{\mathbf{s}}$ $n \times 1$ vector of estimated release function
65	$t$ time
66	$\bar{t}$ sampling time
67	$t_{start}$ starting time
68	$t_{end}$ ending time
69	$\mathbf{u}$ velocity tensor
70	$\mathbf{v}$ $m \times 1$ measurement errors
71	$\mathbf{V}$ $n \times n$ matrix, covariance of the estimate of the errors
72	$\mathbf{x}$ position in the domain
73	$\mathbf{x}_0$ source location
74	$x_{inj}$ longitudinal coordinate of injector
75	$\mathbf{X}_0$ random source location

76	$\mathbf{X}$ $n \times p$ matrix, mean of the unknown process
77	$\mathbf{x}_w$ observation location
78	$\mathbf{z}$ $m \times 1$ observations
79	$z_{inj}$ vertical coordinate of injector
80	$\alpha$ positive number
81	$\alpha_L$ longitudinal dispersivity
82	$\alpha_T$ transversal dispersivity
83	$\beta_x$ normalization factor
84	$\Gamma_i$ boundaries
85	$\delta$ Dirac delta function
86	$\Delta t$ numerical model time step
87	$\Delta x$ longitudinal size of numerical cell grid
88	$\Delta y$ transversal size of numerical cell grid
89	$\Delta z$ vertical size of numerical cell grid
90	$\varepsilon_i$ $i$ th measurement error
91	$\eta$ time
92	$\theta$ structural parameters of the covariance function
93	$\lambda_s$ correlation time length of the unknown release function $s(t)$
94	$\Lambda$ $n \times m$ Kriging coefficients
95	$\Xi$ $m \times m$ dummy matrix
96	$\sigma_R^2$ variance of the measurement error
97	$\sigma_S^2$ variance of the unknown release function $s(t)$
98	$\Sigma$ $m \times m$ dummy matrix
99	$\tau$ backward time
100	$\tau_w$ backward sampling time
101	$\varphi$ porosity
102	$\psi^*$ adjoint state
103	$\nabla$ Nabla operator

## 104 **1 Introduction**

105 Source identification and recovery of the pollutant release history in groundwater have received much attention in  
 106 recent years. The identification of a source location could allow to identify the true cause of the contamination and to  
 107 foresee the future pollution spread, while the release time, the duration, and the maximum value of the released solute  
 108 concentration could allow to apportion remediation costs among the responsible parties.

109 The problems in recovering the release history and/or the source location were studied extensively in the past. In this  
 110 paper two different approaches are compared with respect to the same test case: simultaneous release function and  
 111 source location (in short SRSI) (Butera et al. 2013), and backward location model based on adjoint state method (briefly  
 112 BPM-ASM) (Neupauer and Wilson 1999). Both methodologies were tested on experimental data collected in a  
 113 laboratory sandbox that reproduces an unconfined aquifer (for details see Citarella et al. 2010). This allows to validate

114 the methodologies in a real test case in which all variables are measured and controlled. This is the first time that a well-  
115 known experimental dataset has been used to test and validate these procedures.

116 The first method, developed by Butera et al. (2013), is based on the geostatistical approach proposed by Snodgrass and  
117 Kitanidis (1997) for the one-dimensional uniform flow. Several improvements and applications of the geostatistical  
118 methodology were proposed by Michalak and Kitanidis 2002, 2003, 2004a, 2004b, and by Butera and Tanda (2003) and  
119 Butera et al. (2006, 2013). Given the linearity of the governing differential equation, the approach uses the transfer  
120 functions (TFs) to describe the effect in time, at a certain location of the aquifer, of an impulse release of a pollutant at a  
121 known source. Although TFs can be analytically determined if the problem has a simple flow field, the characteristics  
122 of the groundwater flow field do not allow this solution in many practical applications. To overcome this difficulty, a  
123 numerical procedure to compute TFs was developed by Butera et al. (2004) and applied to homogeneous and weakly  
124 heterogeneous aquifers (Butera et al. 2006). The source identification procedure requires a preliminary delineation of an  
125 area where the pollutant source is most likely to be present, but it allows to obtain the simultaneous identification of the  
126 release history and the source location. An application of this methodology, in a complex real case study, was  
127 implemented by Gzyl et al. (2014).

128 The second methodology was developed by Neupauer and Wilson (1999, 2001). In their works, these authors showed  
129 that the backward location and travel time probability density functions (PDF) are related to adjoint states of  
130 concentration, and they developed a technique for obtaining the governing equation of the backward model using the  
131 adjoint theory. The backward location PDF describes the possible former positions of the observed contamination at a  
132 certain time before the detection, while the backward travel time PDF describes the possible travel time of the  
133 contaminant from a selected upgradient position to the observation location. By using an adjoint model, an  
134 instantaneous point source of an adjoint state (related to the PDFs) is released at the observation location at an observed  
135 time. The adjoint state is thus transported upgradient and backward in time, following the same processes that occur in  
136 forward contaminant transport modeling. The resulting spatial distribution of the concentration in the domain is related  
137 to the backward location PDF. The work by Neupauer and Wilson (1999, 2001) was improved by considering non-  
138 uniform flow field (Neupauer and Wilson 2002), sorbing solutes (Neupauer and Wilson 2004a, 2004b, 2009), multiple  
139 observations (Neupauer and Wilson 2005), measured concentration (Neupauer and Lin 2006), and sorption and decay  
140 phenomena (Neupauer et al. 2007). By using a previously calibrated numerical model, the single-observation and  
141 multiple-observation backward location PDF can be computed. This PDF was then conditioned on the concentration  
142 measurement, by reducing the variance of the backward location PDF and improving the results. However, this  
143 methodology was developed to manage plumes originated from a single instantaneous point source; nevertheless, it  
144 seems an interesting approach and in this work its performance has been tested on a continuous release source.  
145 The manuscript is organized in three parts: first, the mathematical statements concerning the two approaches are  
146 presented, then a brief description of the experimental equipment is reported; finally, the results of the application of the  
147 two methodologies are presented and discussed.

## 148 **2 Mathematical Statements**

### 149 **2.1 Groundwater Transport**

150 Equation (1) describes the transport process in an aquifer corresponding to the injection of a non-sorbing, non-reactive  
151 solute in a point source (Bear and Verruijt 1987):

$$152 \quad \varphi \frac{\partial(C(\mathbf{x}, t))}{\partial t} = \nabla[\varphi \cdot \mathbf{D}(\mathbf{x}) \cdot \nabla C(\mathbf{x}, t)] - \nabla[\varphi \cdot \mathbf{u}(\mathbf{x}, t) \cdot C(\mathbf{x}, t)] + s(\mathbf{x}_0, t) \cdot \delta(\mathbf{x} - \mathbf{x}_0) \quad (1)$$

153 where:  $\varphi$  [-] is the effective porosity (taken as spatially variable, but constant in time),  $\mathbf{u}(\mathbf{x},t)$  [ $L \cdot T^{-1}$ ] is the effective  
 154 velocity at location  $\mathbf{x}$  and time  $t$ ,  $\mathbf{D}(\mathbf{x})$  [ $L^2 T^{-1}$ ] the dispersion tensor,  $C(\mathbf{x},t)$  [ $M \cdot L^{-3}$ ] the concentration at location  $\mathbf{x}$  and  
 155 time  $t$ ,  $s(\mathbf{x}_0,t)=C_0(\mathbf{x}_0,t) \cdot q_0(\mathbf{x}_0,t)$  [ $M \cdot T^{-1}$ ] is the amount of pollutant per time unit injected into the aquifer through the  
 156 source located at  $\mathbf{x}_0$ ,  $C_0(\mathbf{x}_0,t)$  is the concentration injected and  $q_0(\mathbf{x}_0,t)$  [ $L^3 \cdot T^{-1}$ ] is the injection flow rate at  $\mathbf{x}_0$  at time  $t$   
 157 [T].

158 The solution of Equation (1), by considering uniform porosity, when associated with the initial and boundary  
 159 conditions:  $C(\mathbf{x},0) = 0$ ;  $C(\infty,t) = 0$ , is given by the following integral (Jury and Roth 1990):

$$160 \quad C(\mathbf{x},t) = \int_0^t s(\mathbf{x}_0,\eta) f(\mathbf{x},t-\eta) d\eta \quad (2)$$

161 where  $f(\mathbf{x},t-\eta)$  [ $L^{-3}$ ] is the transfer function (TF) that describes the effects at  $\mathbf{x}$  at time  $t$  by an impulse injection occurring  
 162 at  $\mathbf{x}_0$  at time  $\eta$  [T].

## 163 2.2 Geostatistical approach

164 The source position was estimated through the procedure developed by Butera et al. (2013). The method, called  
 165 simultaneous release function and source location identification (SRSI), allows to identify not only the source position  
 166 but also its release history in time.

167 The SRSI procedure can be summarized with the following steps:

- 168 • collect a set of concentration measurements in space and/or time;
- 169 • delineate the suspect area (SA) for the source location and discretize it into  $J$  sub-areas assuming the origin of  
 170 the possible sources in the centroid of any sub-area;
- 171 • compute the transfer functions at the monitoring points for each possible source ( $J$  runs of the numerical  
 172 transport model);
- 173 • recover the release histories performing the geostatistical procedure which simultaneously considers all the

174 possible point sources (superposition effect) by means of  $C(\mathbf{x},t) = \sum_{j=1}^J \int_0^t s_j(\eta) \cdot f_j(\mathbf{x},t-\eta) d\eta$  where  $j$  is one of

175  $J$  generic sub-areas within the SA;

- 176 • identify the source location as the location from which the highest amount of released pollutant is estimated.

177 TFs can be determined easily in simple flow conditions (such as homogeneous and isotropic porous media in absence of  
 178 interferences) by means of analytical formulations (Bear and Verruijt 1987). When considering a non-uniform flow  
 179 field (for instance, heterogeneous porous media, or presence of a withdrawal) it is necessary to develop numerical  
 180 strategies. Butera et al. (2006, 2013) developed the Stepwise Input Function procedure (SIF) for TF calculation. The  
 181 procedure basically consists in making the time derivative of equation (2) by considering a constant and known input  
 182 function  $s(\mathbf{x}_0,t) = F_0 \cdot H_{SF}(t)$ , where  $H_{SF}(t)$  is the Heaviside step function and  $F_0 = C_0 \cdot q_0$  is the amount of pollutant  
 183 injected into the aquifer with constant and known concentration  $C_0$  at flow rate  $q_0$ .

184 The TF results in:

$$185 \quad f(\mathbf{x},t) = \frac{1}{F_0} \frac{\partial C(\mathbf{x},t)}{\partial t} \quad (3)$$

186 Equation (3), coupled with a numerical flow and transport model known from calibration or expert elicitation, allows to  
 187 easily determine the TFs at a generic point  $\mathbf{x}$  by processing the concentration history (breakthrough curve) at the same  
 188 location due to a stepwise tracer injection at  $\mathbf{x}_0$ .

189 The core of the method is the quasilinear geostatistical approach developed by Kitanidis (1995, 1996) and Snodgrass  
 190 and Kitanidis (1997), which is briefly summarized in the following.

191 The observed concentration data at a known time  $\bar{t}$  can be expressed as a function of the release process by the  
 192 following equation:

$$193 \quad \mathbf{z} = \mathbf{h}(\mathbf{s}) + \mathbf{v} \quad (4)$$

194 where  $\mathbf{z}$  is a  $m \times 1$  vector of observations,  $\mathbf{h}(\mathbf{s})$  is the  $n \times 1$  vector containing the time discretization of the unknown release  
 195 function  $s(t)$  and  $\mathbf{v}$  is a  $m \times 1$  vector of epistemic errors with zero mean and known covariance matrix  $\mathbf{R} = \sigma_R^2 \cdot \mathbf{I}$ .

196 In the case of a conservative solute, the relationship between the observed concentration and the release is linear, and  
 197 equation (4) can be simplified to (Snodgrass and Kitanidis, 1997):

$$198 \quad \mathbf{z} = \mathbf{H} \cdot \mathbf{s} + \mathbf{v} \quad (5)$$

199 Equation (5) represents the matrix form of equation (2), where matrix  $\mathbf{H}$  contains the values of the TFs ( $f$ ), computed at  
 200 appropriate times and locations:

$$201 \quad \mathbf{H} = \Delta t \begin{bmatrix} f(\mathbf{x}_1, \bar{t} - \Delta t) & \dots & f(\mathbf{x}_1, \bar{t} - n\Delta t) \\ f(\mathbf{x}_2, \bar{t} - \Delta t) & \dots & f(\mathbf{x}_2, \bar{t} - n\Delta t) \\ \dots & \dots & \dots \\ f(\mathbf{x}_m, \bar{t} - \Delta t) & \dots & f(\mathbf{x}_m, \bar{t} - n\Delta t) \end{bmatrix} \quad (6)$$

202 The transfer matrix  $\mathbf{H}$  includes all the characteristics of the flow and transport processes. Vector  $\mathbf{s}$  can be considered  
 203 random and characterized by an unknown mean  $E[\mathbf{s}] = \mathbf{X}\mathbf{b}$  and a covariance function  $\mathbf{Q}(\boldsymbol{\theta}) = E[(\mathbf{s} - \mathbf{X}\mathbf{b})(\mathbf{s} - \mathbf{X}\mathbf{b})^T]$ , where  
 204  $E[\cdot]$  denotes the expected value,  $\mathbf{X}$  is a  $n \times p$  matrix of known functions,  $\mathbf{b}$  is a vector of size  $p \times 1$  that contains the  
 205 unknown drift coefficients and  $\boldsymbol{\theta}$  are the unknown structural parameters. In this work, a constant but unknown mean is  
 206 considered; thus  $\mathbf{X}$  is an  $n \times 1$  vector filled by 1 and  $\mathbf{b}$  is the scalar unknown mean of the function; moreover a Gaussian  
 207 covariance function has been considered, so  $\boldsymbol{\theta}$  are the variance  $\sigma_s^2$  and the correlation time length  $\lambda_s$ .

208 The estimation procedure proposed by Kitanidis (1995) is divided into two parts: first the structural parameters  $\boldsymbol{\theta}$  of the  
 209 selected covariance function are determined, then the unknown release function is estimated by means of a Kriging  
 210 process. The identification of the structural parameters follows a Restricted Maximum Likelihood approach. The  
 211 probability that the random process with parameter  $\boldsymbol{\theta}$  reproduces observation  $\mathbf{z}$  can be estimated through the following:

$$212 \quad p(\mathbf{z} | \boldsymbol{\theta}) \propto |\Sigma|^{-1/2} \left| \mathbf{X}^T \mathbf{H}^T \Sigma^{-1} \mathbf{H} \mathbf{X} \right|^{-1/2} \exp \left[ -\frac{1}{2} \mathbf{z}^T \Xi^{-1} \mathbf{z} \right] \quad (7)$$

213 where  $\Sigma = \mathbf{H} \mathbf{Q} \mathbf{H}^T + \mathbf{R}$  and  $\Xi = \Sigma^{-1} - \Sigma^{-1} \mathbf{H} \mathbf{X} (\mathbf{X}^T \mathbf{H}^T \Sigma^{-1} \mathbf{H} \mathbf{X})^{-1} \mathbf{X}^T \mathbf{H}^T \Sigma^{-1}$ .

214 Once the structural parameters are computed, by maximizing the probability (Eq. (7)), estimation  $\hat{s}$  of the release  
 215 function  $s(t)$  is obtained through Kriging:

$$216 \quad \hat{s} = \Lambda \cdot \mathbf{z} \quad (8)$$

217 where matrix  $\mathbf{\Lambda}$  ( $n \times m$ ) of the Kriging weights is calculated by solving the following system obtained from the un-  
 218 biasedness and minimum variance conditions:

$$219 \begin{bmatrix} \mathbf{\Sigma} & \mathbf{HX} \\ (\mathbf{HX})^T & \mathbf{0} \end{bmatrix} \begin{bmatrix} \mathbf{\Lambda}^T \\ \mathbf{M} \end{bmatrix} = \begin{bmatrix} \mathbf{HQ} \\ \mathbf{X}^T \end{bmatrix} \quad (9)$$

220 In equation (9),  $\mathbf{M}$  ( $p \times m$ ) is a matrix of Lagrange multipliers. The covariance matrix of the estimation error is:

$$221 \mathbf{V} = -\mathbf{XM} + \mathbf{Q} - \mathbf{QH}^T \mathbf{\Lambda}^T \quad (10)$$

222 A transformation (Box and Cox 1964) of variable  $\mathbf{s}$  was considered in order to enforce non-negativity of the estimated  
 223 concentration; the new unknown becomes:

$$224 \tilde{\mathbf{s}} = \alpha(\mathbf{s}^{1/\alpha} - 1) \quad (11)$$

225 where  $\alpha$  is a positive number and is chosen as small as possible while ensuring  $\tilde{\mathbf{s}} > -\alpha$ .

226 When the values of  $\mathbf{s}$  are constrained to be positive and are physically compatible, equation (4) becomes:

$$227 \mathbf{z} = \mathbf{H} \cdot \mathbf{s} + \mathbf{v} = \mathbf{H} \cdot \left( \frac{\tilde{\mathbf{s}} + \alpha}{\alpha} \right)^\alpha + \mathbf{v} = \mathbf{h}(\tilde{\mathbf{s}}) + \mathbf{v} \quad (12)$$

228 In this case,  $\mathbf{h}(\tilde{\mathbf{s}})$  is not linear with respect to the new unknown  $\tilde{\mathbf{s}}$  and the solution is reached iteratively (for details see  
 229 Kitanidis 1995, 1996).

230 For the SRSI procedure, vector  $\mathbf{s}$  of the unknown release function in (4) is made up by the collection of  $J$  sub-vectors  $\mathbf{s}_j$ ,  
 231 each with dimensions  $n_i \times 1$ , where  $n_i$  is the number of time values used to discretize the release history. The total  
 232 dimension of  $\mathbf{s}$  is:  $(n_1 + n_2 + \dots + n_J) \times 1$ :

$$233 \mathbf{s} = \begin{bmatrix} \mathbf{s}_1 \\ \mathbf{s}_2 \\ \dots \\ \mathbf{s}_J \end{bmatrix} \quad (13)$$

234 The transfer matrix  $\mathbf{H}$  is a block matrix

$$235 \mathbf{H} = [\mathbf{H}_1 \quad \mathbf{H}_2 \quad \dots \quad \mathbf{H}_J] \quad (14)$$

236 whose dimensions are  $m \times (n_1 + n_2 + \dots + n_J)$ . The generic matrix  $\mathbf{H}_j$  describes the effects of the pollutant release in the  
 237 sub-area  $j$  on the concentration data measured at the  $m$  monitoring points. The release history of each source location is  
 238 assumed to be independent of the others. For this reason covariance matrix  $\mathbf{Q}$  of the  $\mathbf{s}$  process is a block matrix with  
 239 non-zero elements in diagonal blocks only.

$$240 \mathbf{Q} = \begin{bmatrix} \mathbf{Q}_1 & 0 & 0 & 0 \\ 0 & \mathbf{Q}_2 & 0 & 0 \\ 0 & 0 & \dots & 0 \\ 0 & 0 & 0 & \mathbf{Q}_J \end{bmatrix} \quad (15)$$

241 The results of the geostatistical procedure described in this section provide the pollutant history in the  $J$  hypothetical  
 242 source locations. The release function in the real source will be substantial, while in the other suspect locations the time  
 243 histories will be negligible.

### 244 2.3 Backward probability model based on adjoint state method (BPM-ASM)



245 The use of backward location probability density function to identify the source location is briefly explained in the  
 246 following. This approach was developed (Neupauer and Wilson 1999) only for an instantaneous point source and it is  
 247 assumed that release time  $\tau$  is known. To calculate backward location PDF a numerical model calibrated on the  
 248 experimental data is often necessary. So, every observation point needs to be considered as an instantaneous source in  
 249 the adjoint equation. The adjoint equations can be solved with the same flow and transport software used in forward  
 250 simulation, on a domain with modified boundary conditions that produce reverse flow direction. In the test case  
 251 explained in the following, one independent backward simulation (flow and transport) for each observation point was  
 252 performed. The initial condition was

$$253 \quad C_w(\mathbf{x}_w, 0) = \frac{1}{Q_w \cdot \Delta t} \quad (16)$$

254 where  $\mathbf{x}_w$  is the position of the observation point, considered now as source,  $Q_w$  [ $L^3T^{-1}$ ] is the instantaneous discharge  
 255 of tracer solution injected, negligible compared to the background flow, and  $\Delta t$  [T] is the time step length. So the mass  
 256 injected in the single time step, is

$$257 \quad M_{injected} = C_w \cdot Q_w \cdot \Delta t = 1 \quad (17)$$

258 The evolution of the plume in the backward model provides the backward location PDF and represents the probability  
 259 that the contaminant exists at that location and given backward time.

### 260 2.3.1 Single-Observation Backward Location PDF

261 For more details see Neupauer and Wilson (1999, 2002, 2003, 2004b). Let us consider the advection dispersion  
 262 equation (Eq. 1) with the following boundary and initial conditions:

$$263 \quad C(\mathbf{x}, t_0) = \frac{m_0}{\varphi} \delta_0(\mathbf{x} - \mathbf{x}_0)$$

$$264 \quad C(\mathbf{x}, t) = g_1(t) \text{ on } \Gamma_1$$

$$265 \quad \left[ \mathbf{D} \frac{\partial C(\mathbf{x}, t)}{\partial x_j} \right] \cdot \mathbf{n} = g_2(t) \text{ on } \Gamma_2 \quad (18)$$

$$266 \quad [\mathbf{u}C(\mathbf{x}, t) - \mathbf{D}\nabla C(\mathbf{x}, t)] \cdot \mathbf{n} = g_3(t) \text{ on } \Gamma_3$$

267 where  $C(\mathbf{x}, t)$  is the concentration,  $\mathbf{D}$  is the dispersion tensor,  $\mathbf{u}$  is the effective velocity,  $t_0$  is the source release time,  $\mathbf{x}_0$  is  
 268 the source location,  $m_0$  [M] is the source mass,  $\delta(\cdot)$  is the Dirac delta function,  $g_i(t)$  are known boundary functions, and  
 269  $\Gamma_i$  are the boundaries.

270 The adjoint of the equation (1) is given by (Neupauer and Wilson 2002)

$$271 \quad \varphi \frac{\partial \psi^*}{\partial \tau} = \nabla(\varphi \mathbf{D} \nabla \psi^*) + \nabla(\varphi \mathbf{u} \psi^*) + q_1 \psi^* \cdot \delta(\mathbf{x} - \mathbf{x}_w) \cdot \delta(\tau - \tau_w) + H_L(\mathbf{x}, \tau) \quad (19)$$

$$272 \quad \psi^*(\mathbf{x}, 0) = 0$$

$$273 \quad \psi^*(\mathbf{x}, \tau) = 0 \text{ on } \Gamma_1$$

$$274 \quad \left[ \mathbf{D} \frac{\partial \psi^*}{\partial x_j} + \mathbf{u} \psi^* \right] \cdot \mathbf{n} = 0 \text{ on } \Gamma_2$$

$$275 \quad \left[ \mathbf{D} \frac{\partial \psi^*}{\partial x_j} \right] \cdot \mathbf{n} = 0 \text{ on } \Gamma_3$$

276 where  $q_I$  [ $T^{-1}$ ] is the source inflow rate per unit volume and  $\psi^*$  [ $L^{-3}$ ] is an adjoint state that is also (Neupauer and  
 277 Wilson 1999) the marginal sensitivity of the concentration to source mass.  $\tau$  is the backward time, while  $\mathbf{x}_w$  and  $\tau_w$  are  
 278 the observation location and backward sampling time, respectively. By using samples taken at monitoring wells

$$279 \quad H_L(\mathbf{x}, \tau) = \left[ \frac{u_1 \alpha_L}{|\mathbf{u}(\mathbf{x})|} \delta'_{x_1}(x_1 - x_{1_w}) \delta(x_2 - x_{2_w}) + \frac{u_2 \alpha_L}{|\mathbf{u}(\mathbf{x})|} \delta'_{x_2}(x_1 - x_{1_w}) \delta(x_2 - x_{2_w}) \right] \delta(\tau - \tau_w) \quad (20)$$

280 where the sample location is  $\mathbf{x}_w = (x_{1_w}, x_{2_w})$ ,  $\delta'_{x_i} = \delta(x_i - x_{i_w})$  is the derivative of the Dirac delta function with respect to  
 281  $x_i$ , vertical bars denote magnitude, and  $\alpha_L$  [ $L$ ] is the longitudinal dispersivity.

282 The relationship between the adjoint state from (19) and the backward location probability density function is given by  
 283 (Neupauer and Wilson 2002)

$$284 \quad f_{\mathbf{x}}(\mathbf{x}; \tau, \mathbf{x}_w, \tau_w) = \varphi \cdot \psi^*(\mathbf{x}; \tau, \mathbf{x}_w, \tau_w) \quad (21)$$

285 where  $f_{\mathbf{x}}(\mathbf{x}; \tau, \mathbf{x}_w, \tau_w)$  is the backward location PDF concerning position  $\mathbf{x}$  at backward time  $\tau$  of a contaminant particle  
 286 that was observed at  $\mathbf{x}_w$  at backward time  $\tau_w$ , and  $\psi^*$  is the adjoint state obtained from (19).

### 287 2.3.2 Multiple Observations Backward Location PDF

288 If several observations are available, multiple-observation backward location density function (Neupauer and Wilson  
 289 2005) can be obtained by first calculating the single-observation PDF for each observation, then combining them. Each  
 290 observation gives additional information to characterize the former position of contamination, thereby reducing the  
 291 uncertainty or variance of the backward location PDF. In this step only the presence of the pollutant at monitoring  
 292 points is considered, and not the concentration values. Let  $N$  be the number of observations, and let  $\{\mathbf{x}_w\}$  and  $\{\tau_w\}$  be  $N$ -  
 293 length vectors of sampling locations and backward sampling times, respectively. The multiple-observation backward  
 294 location PDF,  $f_{\mathbf{X}}(\mathbf{x}; \tau, \{\mathbf{x}_w\}, \{\tau_w\})$ , describes the possible former positions of all contaminant particles observed, given  
 295 that at a previous time  $\tau$ , they were at the same location, i.e. source location. It is calculated as

$$296 \quad f_{\mathbf{X}}(\mathbf{x}; \tau, \{\mathbf{x}_w\}, \{\tau_w\}) = \frac{\prod_{i=1}^N f_{\mathbf{X}}(\mathbf{x}; \tau, \mathbf{x}_{wi}, \tau_{wi})}{\int \prod_{i=1}^N f_{\mathbf{X}}(\mathbf{x}; \tau, \mathbf{x}_{wi}, \tau_{wi}) d\mathbf{x}} \quad (22)$$

297 where  $f_{\mathbf{X}}(\mathbf{x}; \tau, \mathbf{x}_{wi}, \tau_{wi})$  is the backward location PDF for the  $i$ th sample, obtained from (21). The sum of multiple-  
 298 observation backward location PDF, calculated on the whole domain, is equal to 1. An important aspect of this method  
 299 is that only the presence or the absence of the pollutant is considered and the concentration values are not taken into  
 300 account.

### 301 2.3.3 Conditioned Backward Location PDF

302 Neupauer and Lin (2006) improved the backward location probability density function conditioning on measured  
 303 concentrations collected after an instantaneous release originated from a point source located in  $\mathbf{x}_0$ .

304 In this case the source mass and the source location are unknown but the measured concentrations are known. Neupauer  
 305 and Lin (2006) suggested using Bayes' theorem with the aim of constraining source mass  $M$  and random source  
 306 location  $\mathbf{x}_0$  on the specific concentrations measured.

307 Let  $\hat{C}_i = \hat{C}(\mathbf{x}_{wi}, \tau_{wi})$ ,  $i = 1, 2, \dots, N$ , the observed concentrations, where  $\mathbf{x}_{wi}$  is the location and  $\tau_{wi}$  is the backward time  
 308 at which sample  $i$  was taken.  $\hat{C}_i$  is considered as a random variable with true value  $C(x_{wi}, \tau_{wi} | m_0, \mathbf{x}_0, \tau)$ . So,  $\varepsilon_i$  is the  
 309 error of the  $i$ th measurement and it is normally-distributed with zero mean and variance  $\sigma_i^2$ . Therefore measured

310 concentrations,  $\hat{C}_i$  are normally distributed with a mean equal to the true concentration and variance  $\sigma_R^2$ , and its PDF is  
 311 given by

$$312 \quad f_{\hat{C}_i}(\hat{C}_i | m_0, x_0, \tau) = \frac{1}{\sqrt{2\pi\sigma_R^2}} \exp\left\{-\frac{[\hat{C}_i - C(x_{wi}, \tau_{wi} | m_0, x_0, \tau)]^2}{2\sigma_R^2}\right\} \quad (23)$$

313 If  $\varepsilon_i$  is considered independent from any other, for a known source mass  $m_0$  and a source location  $\mathbf{x}_0$ , the joint PDF on  
 314 all  $N$  measured concentrations is simply the product of the PDFs for the individual observations given by

$$315 \quad f_{\hat{\mathbf{C}}}(\hat{\mathbf{C}} | m_0, \mathbf{x}_0, \tau) = \prod_{i=1}^N \frac{1}{\sqrt{2\pi\sigma_R^2}} \exp\left\{-\frac{[\hat{C}_i - C(x_{wi}, \tau_{wi} | m_0, x_0, \tau)]^2}{2\sigma_R^2}\right\} \quad (24)$$

316 Assuming that the source mass is independent of source location in the absence of any concentration information, the  
 317 final result is (for more details see Neupauer and Lin 2006)

$$318 \quad f_{\mathbf{x}_0 | \hat{\mathbf{C}}}(\mathbf{x}_0 | \hat{\mathbf{C}}; \tau) = \beta_x \int \prod_{i=1}^N \left[ f_{\hat{C}_i | M, X_0}(\hat{C}_i | m_0, \mathbf{x}_0; \tau) \cdot f_x(\mathbf{x}; \tau, \mathbf{x}_{wi}, \tau_{wi}) \right] \cdot dm_0 \quad (25)$$

319 where  $f_{\hat{C}_i | M, X_0}(\hat{C}_i | m_0, \mathbf{x}_0; \tau)$  is the distribution of  $\hat{C}_i$ , calculated for each observation for a range of possible source  
 320 masses and a range of possible source locations that includes the entire spatial domain of the model,  $f_x(\mathbf{x}; \tau, \mathbf{x}_{wi}, \tau_{wi})$  is  
 321 the unconditioned backward location probability density function and  $\beta_x$  is a normalization factor.

## 322 **3 Test case**

### 323 **3.1 Description of the experimental equipment**

324 The transport experiments were performed in a laboratory device (sandbox) built with polymethyl methacrylate  
 325 (PMMA) plates. The sandbox reproduces an unconfined aquifer governed by two levels (upstream and downstream).  
 326 The external dimensions of the sandbox are 1.20m  $\times$  0.14m  $\times$  0.73m. Along the longest axis  $x$ , the sandbox is made of  
 327 three parts (Fig. 1): two tanks (upstream and downstream), which allow the regulation of the water level and, as a  
 328 consequence, of the flux, and a central chamber (0.95 m  $\times$  0.10m  $\times$  0.70m) which contains the porous medium. The  
 329 water discharge is monitored with a flow meter. The porous medium consists of glass beads with diameter in the range  
 330 between 0.75 and 1 mm; the hydraulic parameters are summarized in Table 1. An injector was positioned in the  
 331 upstream part of the sandbox (see Fig. 1), and fluorescein sodium salt was chosen as tracer because, when mixed with  
 332 water and excited with blue light ( $\lambda = 490$  nm), it irradiates in longer wavelength (green light,  $\lambda = 520$  nm). The  
 333 experimental device was placed in a darkroom to avoid all external light contamination and lightened by 8  
 334 monochromatic blue LEDs. All variables, such as upstream and downstream level, injected discharge, temperature,  
 335 background discharge, start and ending of injection, were acquired by means of a data acquisition system. The  
 336 luminosity at each point of the sandbox was recorded by a digital camera and then converted in concentration through  
 337 an imaging technique. The mass released by the injector and the one estimated through the image processing were  
 338 compared to evaluate the reliability of data collected. Another confirmation of the validity of the data used was obtained  
 339 by comparing the mass rate that flows through the sandbox with the known injected one. During the device calibration,  
 340 the maximum measurement error of concentration  $\sigma_R$  was estimated as less than 3 mg/l. A detailed description of the  
 341 sandbox can be found in (Citarella et al. 2010).

### 342 **3.2 Description of the test**

343 A tracer solution with variable mass rate was injected by using the experimental device described in the previous  
344 section. The average background flow rate was measured as 25 ml/s, obtained with upstream and downstream constant  
345 heads of 59.9 cm and 53.6 cm, respectively. The injector was located at the coordinates  $x_{inj} = 14.25$  cm,  $z_{inj} = 32.75$  cm,  
346 and it was as wide as the central chamber.

347 The test had a time length of 2,200 s. The injection started at time  $t_{start} = 310$  s and finished at  $t_{end} = 1,800$  s, so it had a  
348 duration of 1,490 s. The mass rate varied in a range between 0 and approximately 60  $\mu\text{g/s}$ . Through a photographic  
349 survey, the tracer concentration was estimated at every point of the domain. Two observation points only were  
350 considered for the application of the two methodologies discussed in this paper (Figs. 3 and 4). At such points the  
351 concentration was estimated every 5 s for the whole duration of the test (Fig. 4). The contaminant release history, in  
352 terms of mass rate, presented three peaks of different magnitude (Fig. 5). In Figs. 4 and 5 the time scale is referred at  
353  $t_{start}$ . Two monitoring points only were considered so as to simulate a real case, in which there are generally few points  
354 and more observations in time, only. Indeed, the realization of a large number of monitoring wells can be very  
355 expensive; moreover, a case with a small number of observation points is more severe.

### 356 **3.3 Description of the numerical model**

357 The sandbox was represented by a numerical flow and transport model in order to provide an essential tool for the two  
358 methodologies.

359 The main assumption is that the flow and transport phenomena are uncoupled and the flow has mainly two-dimensional  
360 components in the vertical plane; the porous medium was considered homogeneous and isotropic. The groundwater  
361 flow was reproduced with MODFLOW 2000 (Harbaugh et al. 2000) and the transport process using MT3DMS with  
362 TVD as advection solver package (Zheng and Wang 1999).

363 The sandbox was described by using a finite difference grid of 192 columns (representing the longitudinal size), 140  
364 layers (describing the sandbox height) and one row only (the depth size), resulting in 26,880 computation nodes. The  
365 size of each numerical cell was  $(\Delta x, \Delta y, \Delta z) = (0.5, 10.0, 0.5)$  cm. The hydraulic parameters (see Table 1) of the  
366 numerical flow model were set up by comparing the measured and computed flows in steady state condition (without  
367 any injection), assuming that the porous medium is homogeneous and isotropic. The transport parameters (see Table 1)  
368 were calibrated by fitting the experimental and numerical breakthrough curves of different experiments (with various  
369 injection rates and concentration values) at several monitoring points. As an example, Figure 6 shows the good  
370 agreement between the numerical and the observed breakthrough curves at three monitoring points located at different  
371 distances (14.75, 33.75, 57.75 cm) downstream the injection point of the present test case.

372 Once the numerical model was validated, it was used for both inverse methodologies: to compute the TFs numerically  
373 (for more details see Butera et al 2013), and to build the backward probability model (Neupauer and Wilson 1999).

### 374 **3.4 Results of the geostatistical approach (GA)**

375 The SA was assumed upgradient from the measurement points in the region  $11.25 \leq x \leq 17.25$  cm and  
376  $29.75 \leq z \leq 35.75$  cm, and it was subdivided into 9 sub areas as shown in Figure 3. The centroid of each area represents  
377 a possible source location. The suspected area is a 6x6 cm square, which is quite large compared to the pore dimension  
378 (1 mm) and the dispersivity values (1.6 mm). In fact it is about 60 times the grain dimension and it is comparable with  
379 the plume's transverse dimension. At first, the TFs relevant to the two monitoring points and the nine possible sources  
380 were computed by applying the SIF procedure, requiring nine runs of the forward flow and transport models. In the test  
381 case considered the mass rate released in time and source position are the unknowns. For each monitoring point, 32  
382 concentration values were considered available throughout the duration of the injection with a time step equal to 70 s

383 (Fig. 4). Uncertainty associated with each concentration value was considered constant as described in section 3.1. A  
384 Gaussian covariance function was chosen  $Q(t_i, t_j | \theta) = \sigma_s^2 \cdot \exp\left(-|t_i - t_j| / \lambda_s^2\right)$ , so variance  $\sigma_s^2$  and correlation length  
385  $\lambda_s^2$  were estimated for each of the nine suspected point source. Finally, the SRSI procedure was carried out and the  
386 release function for each of the nine possible sources was obtained (Fig. 5).

387 The release history assumes negligible values in all the locations, except in  $x = 14.25$  cm,  $z = 32.75$  cm. This result  
388 indicates that the source is located in the sub-area with those centroid coordinates (just the ones of the actual source).  
389 The release history recovered for the source location  $x = 14.25$  cm,  $z = 32.75$  cm shows a good agreement between the  
390 observed and the peak times recovered; moreover the whole release history is included in the 5-95% confidence  
391 interval.

392 A drawback of methodology is that if the number of SA is too high, the methodology could be computationally  
393 ineffective, because it requires one run of the numerical model for each suspected source.

### 394 **3.5 Results of backward model based on adjoint state method (BPM -ASM)**

#### 395 **3.5.1 Unconditioned single-observation backward location PDF**

396 At first, the single unconditioned backward location probability density functions were calculated at the single  
397 observation points P1 and P2 (see Fig. 3). At backward time  $\tau = 0$ , a tracer solution discharge equal to  $Q_w$  and a  
398 concentration given by (16), was injected for one time step only.

399 The travel times between the true source and observation points P1 and P2 are respectively  $\Delta\tau_{P1} = 340$  s and  
400  $\Delta\tau_{P2} = 400$  s. So, the evolution of the backward location PDF, at backward time  $\tau = 400$  s (Figs. 7a and 7b) was  
401 computed. Whereas the porous medium is homogeneous and isotropic, the shapes of the contours are symmetric as  
402 shown in Figs. 7a and 7b. Moreover, Fig. 7a shows that for a closer observation point, the backward location PDF is  
403 higher, as expected. This spreading out of probability density function is caused by dispersion, similar to the processes  
404 that occur in a forward model.

#### 405 **3.5.2 Multiple-observation backward location PDF**

406 In this case, two observations of contamination collected at the two observation points are used to calculate the  
407 multiple-observation location PDF, given by (22). It was assumed that the particles observed in P1 and P2 were  
408 originated by a unique release at the same point source location at the same backward time. As shown in Figure 7c, the  
409 dispersion of the backward location is smaller than the two single-observation location PDF (Figs. 7a and 7b). This  
410 approach does not allow to consider the concentration measurement, but only the presence or the absence of the  
411 contaminant. So, one point detected in a marginal position of the plume could compromise the result because the  
412 observations cannot identify the plume shape. In fact, in the backward model, the backward location PDF simply  
413 follows the same processes that occur in the forward model, and the resulting multiple-observation backward location  
414 PDF is not able to give a weight to observations with a higher concentration value.

#### 415 **3.5.3 Backward location PDF conditioned on measured concentrations**

416 In the previous sections the concentration measured at the two observation points was not considered. A lot of  
417 information about the concentration in time at both points P1 and P2 was available through the images collected during  
418 the laboratory test. However the method requires only the concentration values at the monitoring points collected at a  
419 time  $\bar{t}$ . Several sample times were considered during preliminary studies and all yielded very similar results. In this  
420 case, the results due to the concentration sampled at  $\bar{t} = 590$  s after  $t_{start}$  are shown. Since the release mass is considered

421 unknown, as the source location,  $f_{\hat{C}_i|M,x_0}(\hat{C}_i|m_0, \mathbf{x}_0; \tau)$  was calculated for each observation for a range of possible  
422 source masses ( $50 \text{ mg} < m_0 < 70 \text{ mg}$ ) and a range of possible source locations that include the entire spatial domain. An  
423 iterative procedure was developed with Matlab (Mathworks 2010) in order to change the source position in every cell,  
424 and run the flow and transport models. This approach could be computationally challenging if the number of cells is too  
425 high, due either to an excessively extensive investigated area or a high spatial resolution. This could be improved by  
426 analyzing the delineated suspected area only, as in the SRSI method. The concentrations observed at the monitoring  
427 points used to condition the backward location PDF were 16.97 mg/l for P1 and 3.88 mg/l for P2 (Fig. 4). As described  
428 in section 3.1, the measurement error is taken as being constant and equal to 3 mg/l, although this value represents the  
429 maximum error, and for small concentrations it is smaller than the one considered. The source location, based on the  
430 backward location probability density function, conditioned on concentrations measured, is very well identified. The  
431 conditioned backward probability PDF plume has a long and narrow shape, as shown in Figure 7d. This improvement  
432 reduces the backward location variance and it allows to increase information on the source.

### 433 **3.5.4 Sensitivity of BPM-ASM on the assumed release time**

434 An assumption on the release time is required when the backward probability model is used, so a study about the effect  
435 of this parameter on the results was considered. In the previous sections, the backward release time was fixed at  $\tau =$   
436 400 s. An error of 10% and 20% of  $\Delta\tau_{P2}$  was considered with the aim of studying the results with a wrong time  
437 release. Thus, backward release times  $\tau_{-10\%} = 360$  s, and  $\tau_{-20\%} = 320$  s were hypothesized, and the results are shown in  
438 Figures 8 and 9. While the results obtained with a 10% error remain acceptable, a 20% error does not allow to identify  
439 the true source location. Moreover, by considering an error of +10% ( $\tau_{+10\%} = 440$  s) or +20% ( $\tau_{+20\%} = 480$  s), the source  
440 location will be estimated upstream the true one and it will present a larger backward location PDF plume than the one  
441 estimated with -10% or -20%.

## 442 **4. Discussion and Conclusions**

443 A comparison between two methods, simultaneous release function and source location identification (SRSI) and  
444 backward probability model based on adjoint state method (BPM-ASM), was carried out. For the application of these  
445 methodologies, experimental data collected through laboratory equipment under controlled conditions were used. This  
446 allowed to perform both methodologies in a real test case in which the true solution was known a priori. A numerical  
447 model was calibrated on experimental data with the aim of computing the transfer functions (Butera et al. 2004, 2006)  
448 and building the backward model for the BPM-ASM (Neupauer and Wilson 1999, 2000). The SRSI procedure is able to  
449 simultaneously recover the release function and identify the source location, but it requires a preliminary delineation of  
450 the suspected areas and some weak hypothesis about the statistical structure of the unknown release function. The  
451 number of preliminary runs of the forward transport model needed to obtain the numerical TFs is equal to the number of  
452 the suspect source locations.

453 The backward location PDF describes the possible former positions of the observed contamination at a fixed time in the  
454 past. By using an adjoint model, an instantaneous point source of an adjoint state (related to the backward location  
455 PDFs) is released at the observation location. Given the simplicity of the flow field, the exchange of the boundary  
456 conditions was sufficient to obtain the backward model. At first, the unconditioned backward-location PDFs were  
457 calculated for each monitoring point, by considering the known time release. The multiple-observation PDF was then  
458 calculated and finally it was conditioned on the experimental concentration measurement. This method identifies the  
459 true source very well. As with the SRSI method, two observation points were used, but while in the SRSI multiple times

460 were considered (32 sampling times for each monitoring point in this test case), in the BPM-ASM only two  
461 observations at a specific time, were required.

462 In conclusion, both methods allow to identify the true source; the SRSI also allows to recover the release history. While  
463 the SRSI method requires the definition of a suspected area, the BPM-ASM requires a known release time: both  
464 hypotheses could be strong in certain cases. Finally both methodologies could be computationally inefficient, and if the  
465 SRSI requires a forward run for each suspected source, the conditioning on the measured concentrations of multiple-  
466 observation PDF requires a forward run for each grid node of the numerical model. The number of the forward runs in  
467 the BPM-ASM could be reduced by analyzing the suspected area only. Another important difference between the two  
468 methods is that the SRSI works with multiple sources too, while the BPM-ASM is able to recognize one point source  
469 only. In fact the multiple-observation PDF is based on the hypothesis that all particles observed at a specific time were  
470 originated at the same location from an instantaneous release. It is important to notice that the BPM-ASM performed  
471 well in the test case considered too, in which the observations were originated from a complex release history.

472 In order to overcome the limits of the two methods, a new procedure which includes the best performance of both  
473 approaches could be developed: for instance the BPM-ASM could be preliminarily used to identify the suspected areas,  
474 by considering several backward times, and then the SRSI can be applied to estimate the true source location and its  
475 release function.

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547

548

## Tables

549

550 **Table 1** Transport and hydraulic parameters of the numerical model.

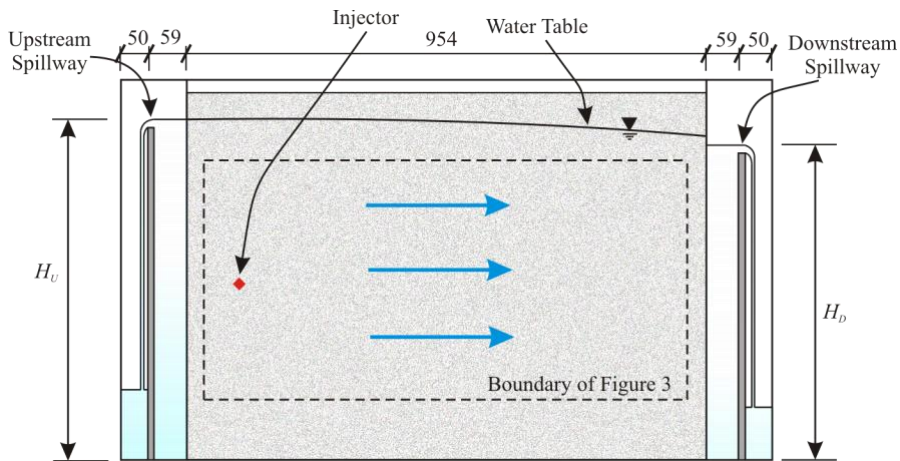
Hydraulic Conductivity [cm/s]	0.652
Porosity	0.37
Specific Storage Coefficient [cm <sup>-1</sup> ]	10 <sup>-4</sup>
Longitudinal Dispersivity [cm]	0.16
Transverse Dispersivity [cm]	0.05

551

552

## Figures

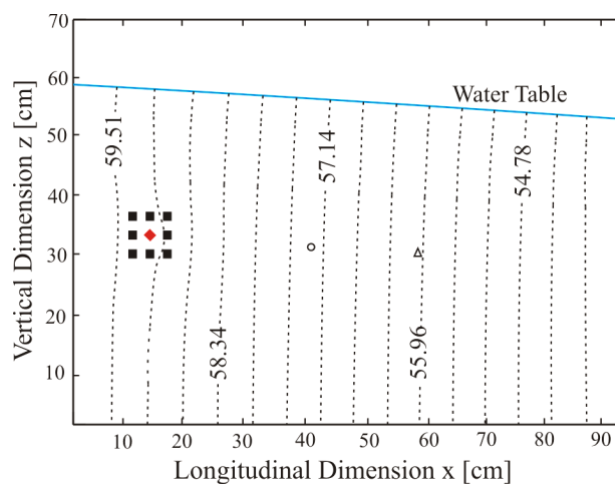
553



554

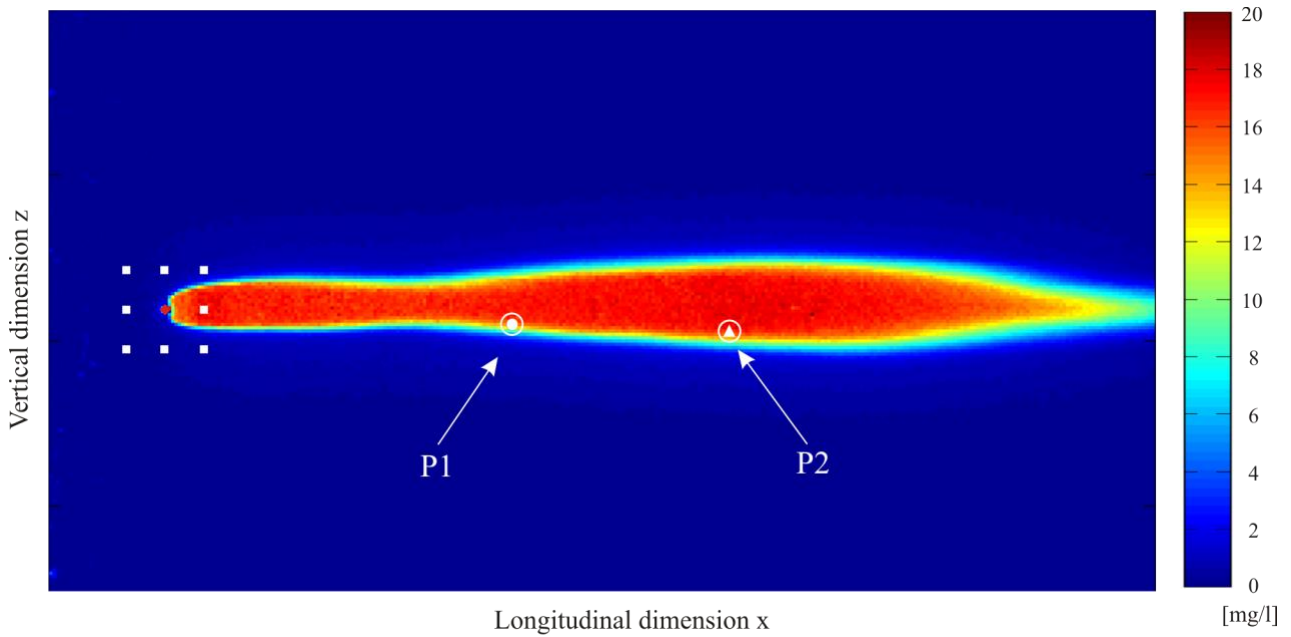
555 **Figure 1** Sketch of the experimental device (lateral view): constant head boundaries upstream ( $H_U$ ) and downstream  
556 ( $H_D$ ); the red diamond is the source location. Dimensions are in mm.

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559 **Figure 2** Hydraulic head distribution (in cm) from the numerical model. The red diamond is the source location. The  
560 filled squares denote the possible sources in the SA. The open circle and the triangle represent observation points P1  
561 and P2, respectively. The experimental equipment reproduced an unconfined aquifer.



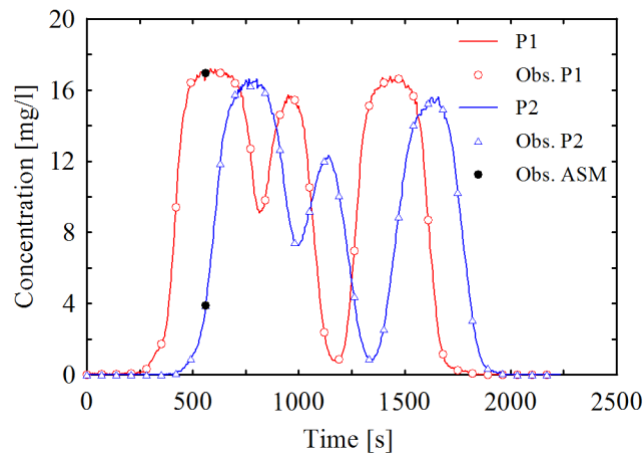
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563 **Figure 3** Concentration field estimated through the analysis of the image collected 690 s after the start of the injection.

564 The concentration field depicted (85.28 x 44.15 cm) corresponds to the rectangle indicated in Figure 1. The red

565 diamond is the source location. The white filled squares denote the possible sources in the SA.

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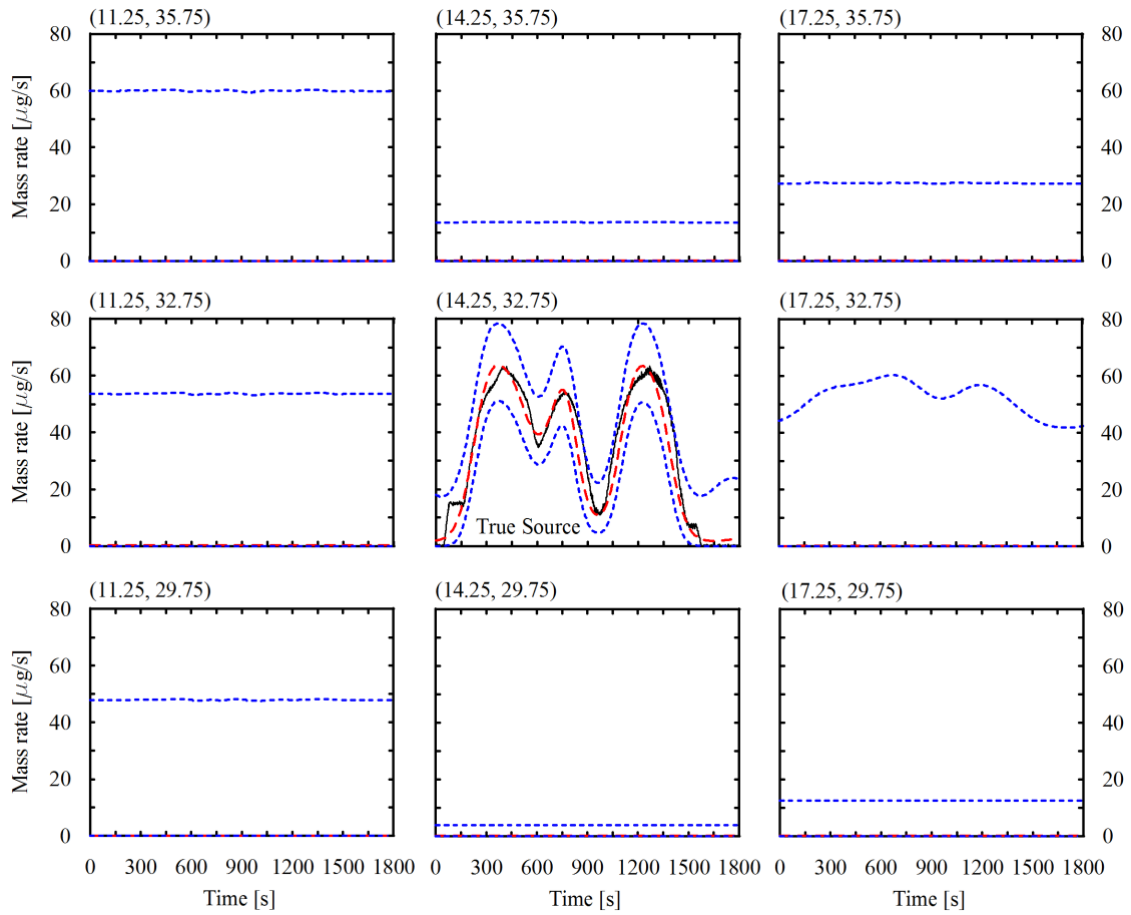


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568 **Figure 4** Concentration observed at the two monitoring points (solid line). Black dots are the observations used to

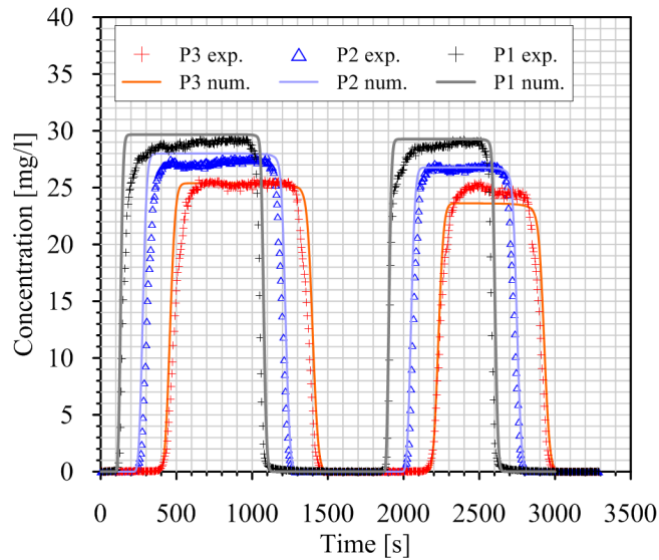
569 condition the multiple-observation PDF in BPM-ASM method. Time 0 s represents the time at which injection starts

570 ( $t_{start}$ ).



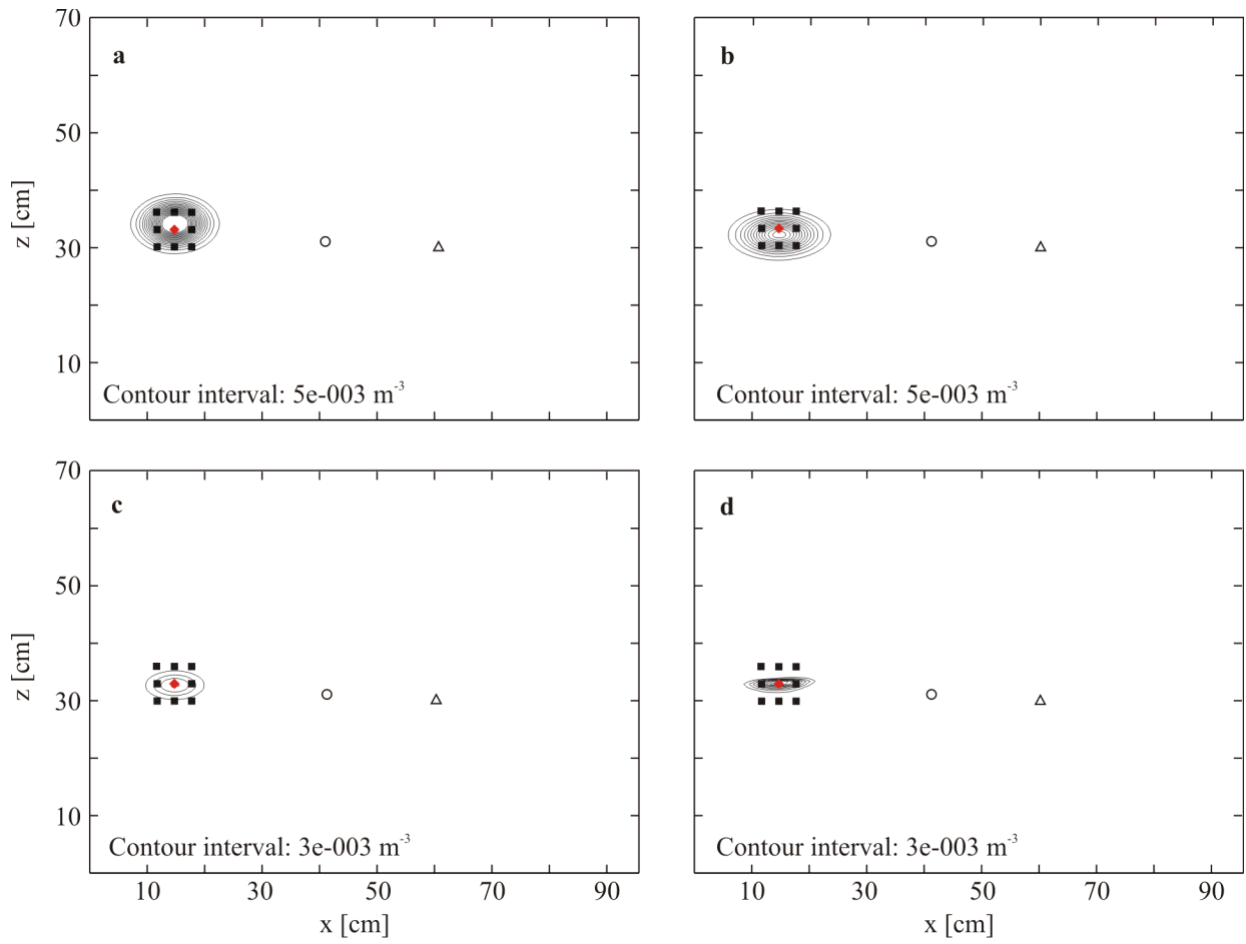
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**Figure 5** The release history recovered at the hypothesized source locations: the true solution (solid black line), best estimate (red dashed line) and 5-95 % confidence interval (blue dotted line). Coordinates of the sources in cm. Time 0 s represent the time at which injection starts ( $t_{start}$ ).



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**Figure 6.** Comparison between numerical and experimental breakthrough curves.



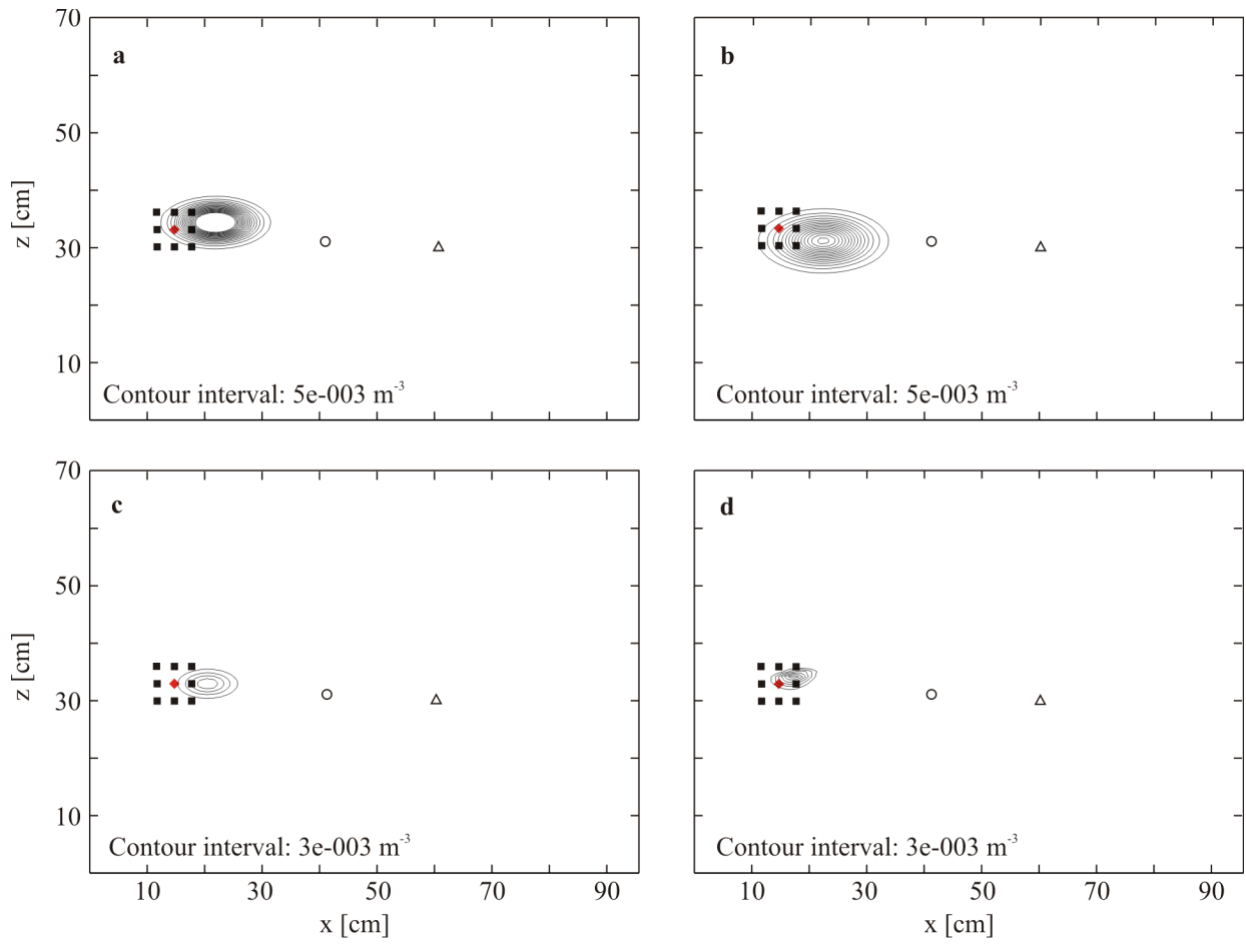
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**Figure 7** Backward location probability density function at  $\tau = 400$  s. (a) Unconditioned backward location PDF: observation point P1. (b) Unconditioned backward location PDF: observation point P2. (c) Multiple-observation (P1 and P2) backward location PDF (d) Conditioned Backward Location PDF on concentrations.



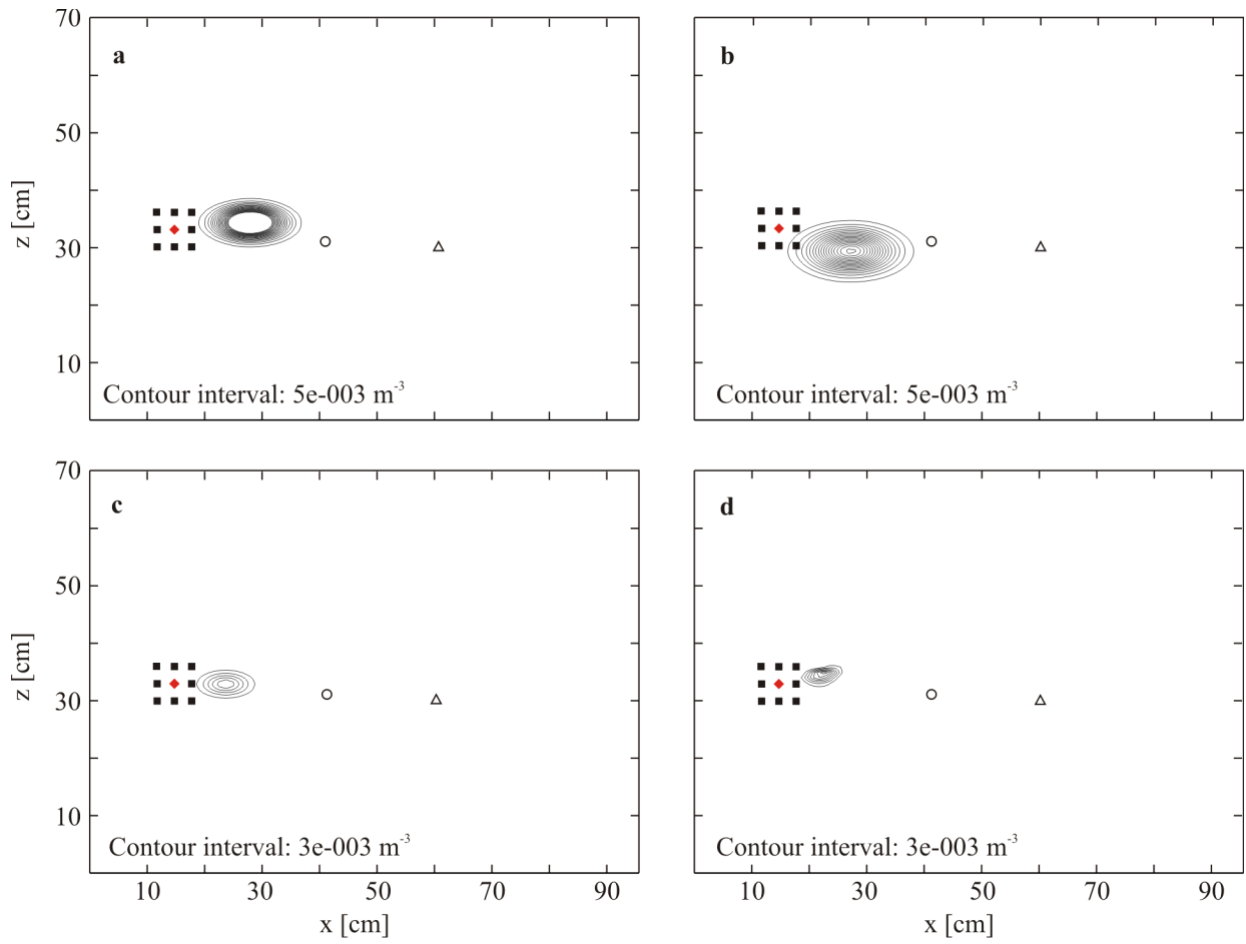
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**Figure 8** Backward location probability density function at  $\tau_{-10\%} = 360$  s. (a) Unconditioned backward location PDF: observation point P1. (b) Unconditioned backward location PDF: observation point P2. (c) Multiple-observation (P1 and P2) backward location PDF (d) Conditioned Backward Location PDF on concentrations.



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**Figure 9** Backward location probability density function at  $\tau_{-20\%} = 320$  s. (a) Unconditioned backward location PDF: observation point P1. (b) Unconditioned backward location PDF: observation point P2. (c) Multiple-observation (P1 and P2) backward location PDF (d) Conditioned Backward Location PDF on concentrations.